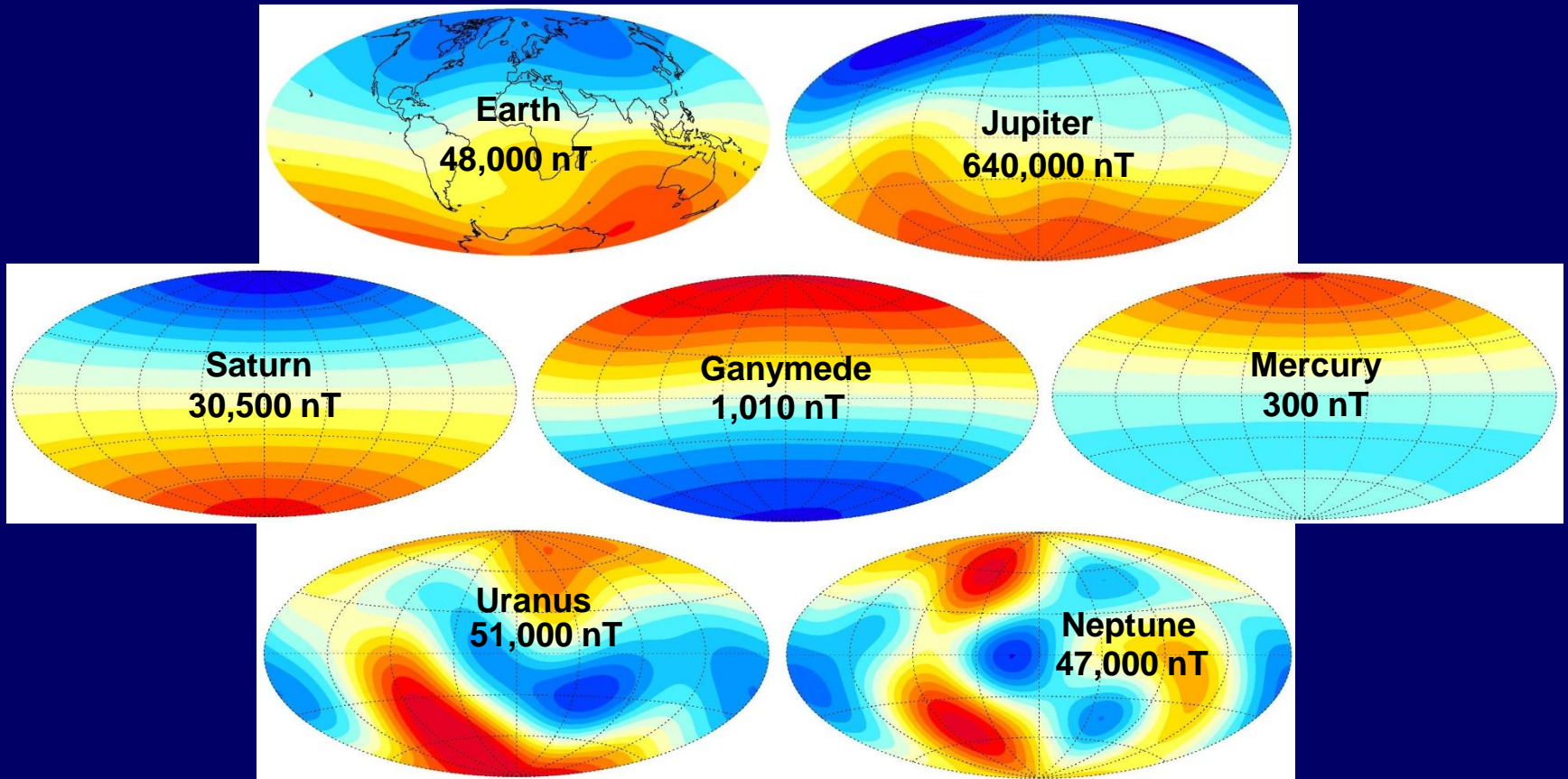
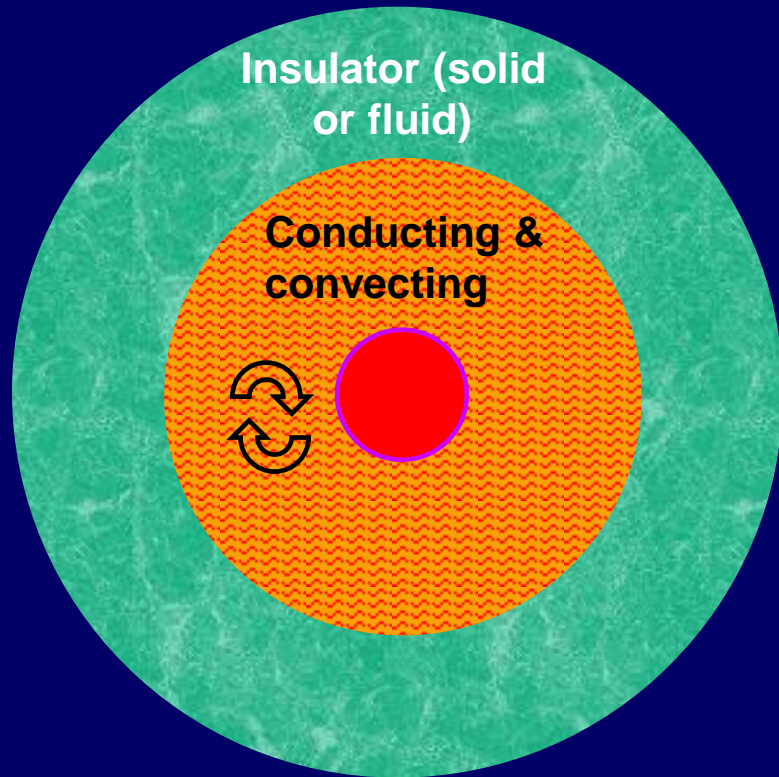


# Planetary dynamos below stably stratified layers

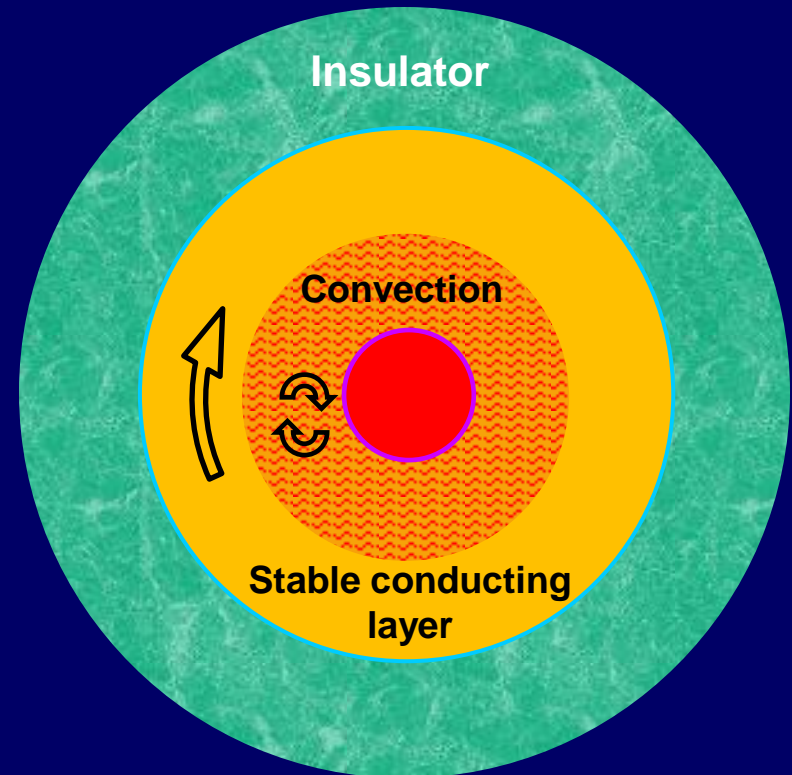
Ulrich Christensen, MPS, Göttingen, Germany



# Standard model



# Stable layer



## Causes for stable layer

Sub-adiabatic T-gradient in outer parts of core

Compositional stratification (e.g. associated with phase separation)

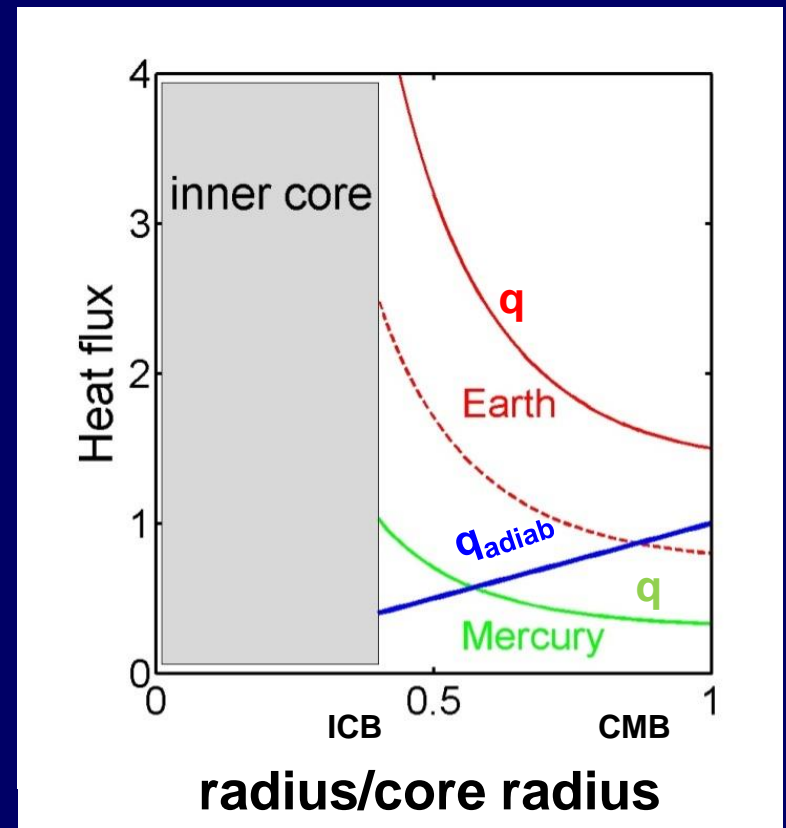
# Thermal convection in cores of terrestrial planets

- Large heat flux  $q_{\text{adiab}}$  can be conducted along adiabatic T-gradient
- Main heat source latent heat of freezing of solid inner core:

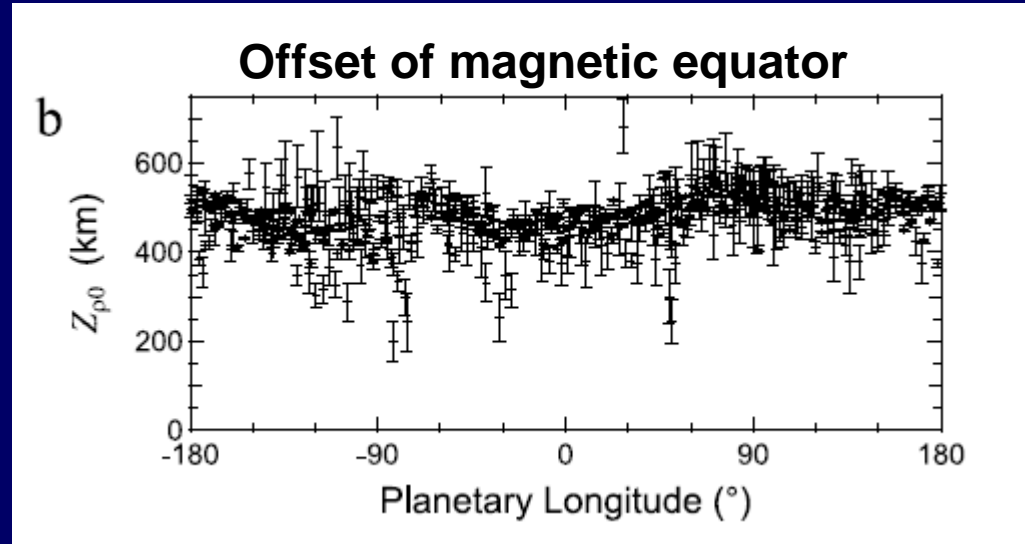
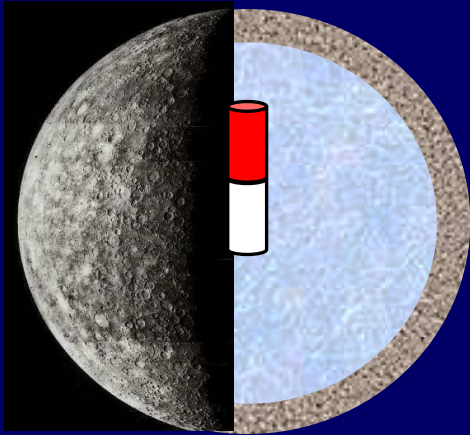
$$q \propto 1/r^2$$

$$q_{\text{adiab}} \propto r$$

- Possible scenario: thermally unstable at depth, stable near core-mantle boundary



# Mercury's magnetic field

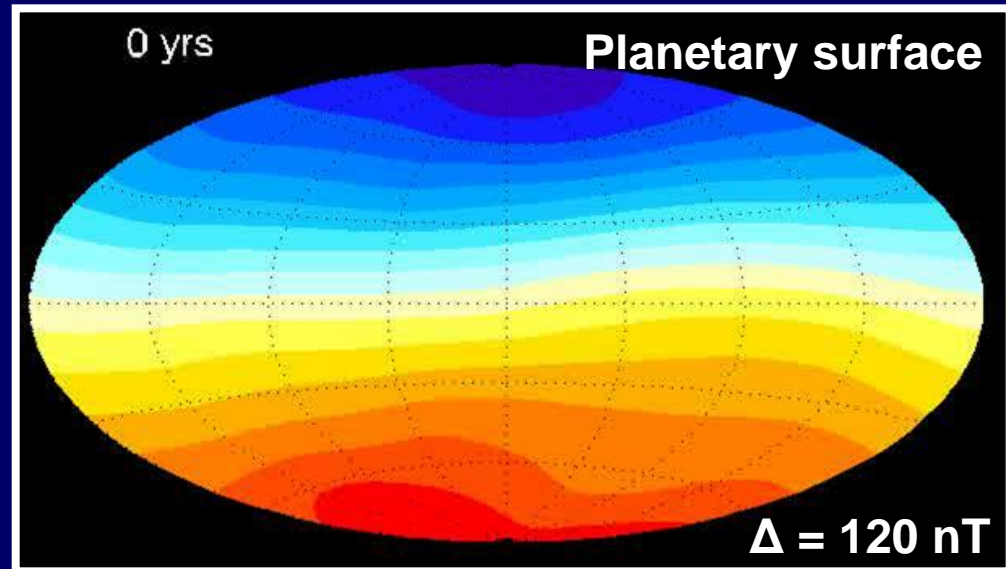
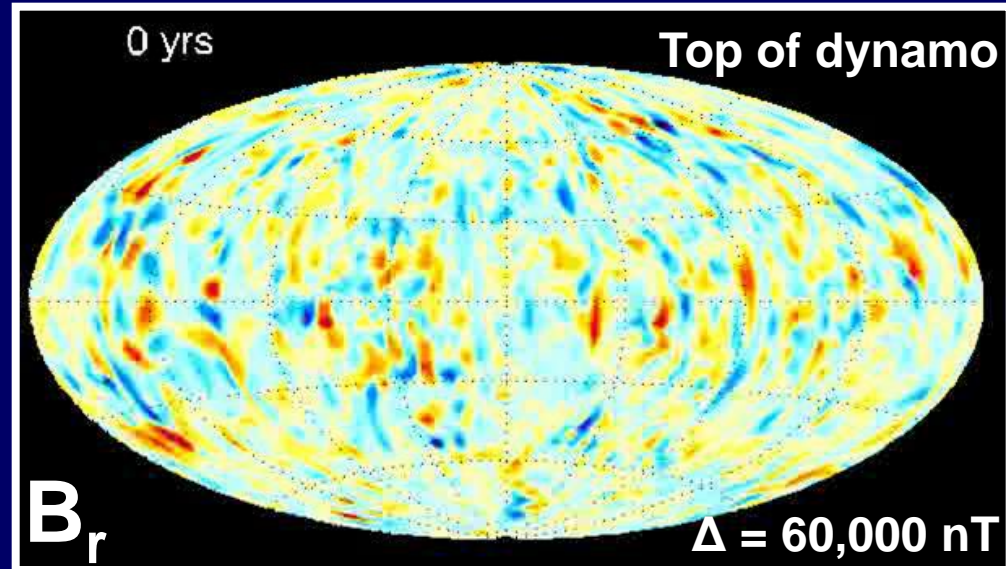
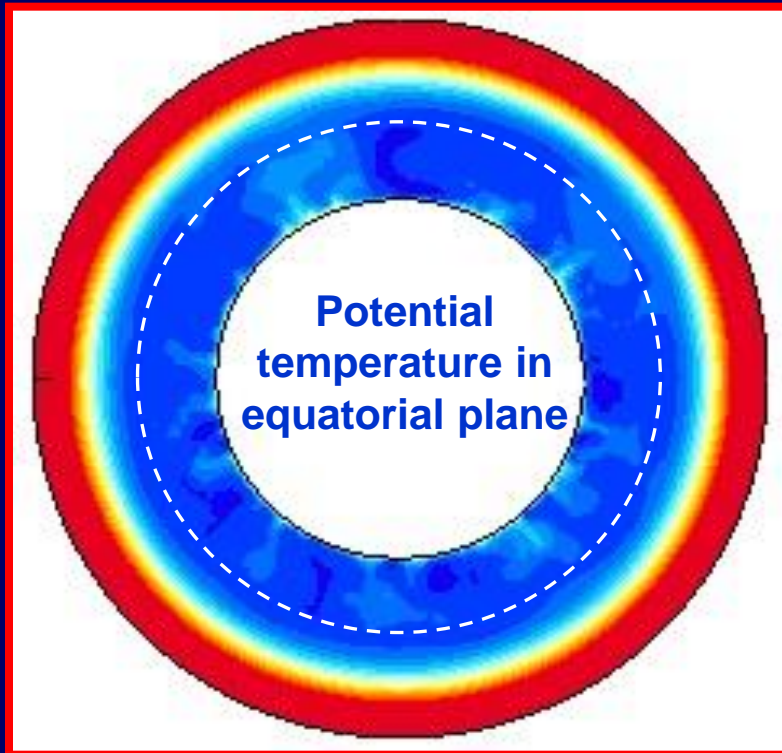


- **Mercury is slow rotator (59 days)**
- **Field dipole-dominated but weak ( $g_{10}=190$  nT)**
- **Dipole tilt wrt rotation axis small ( $< 0.8^{\circ}$ )**
- **Relatively large axial quadrupole ( $g_{20}/g_{10} = 0.39$ )**

(Anderson et al., 2011, 2012)



# Mercury dynamo below stable layer

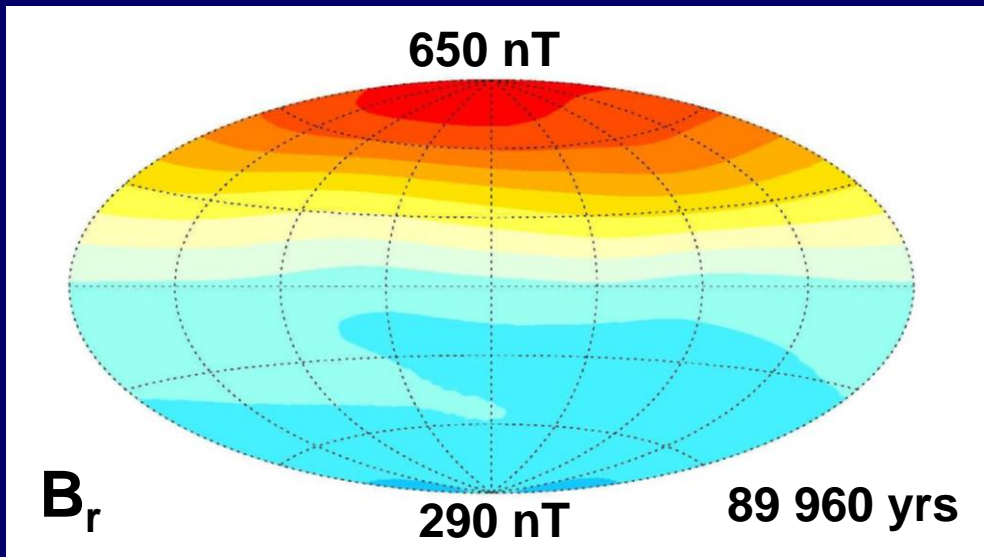


- Internal field strong, small-scale
- Surface field weak, large-scale

Christensen, 2006

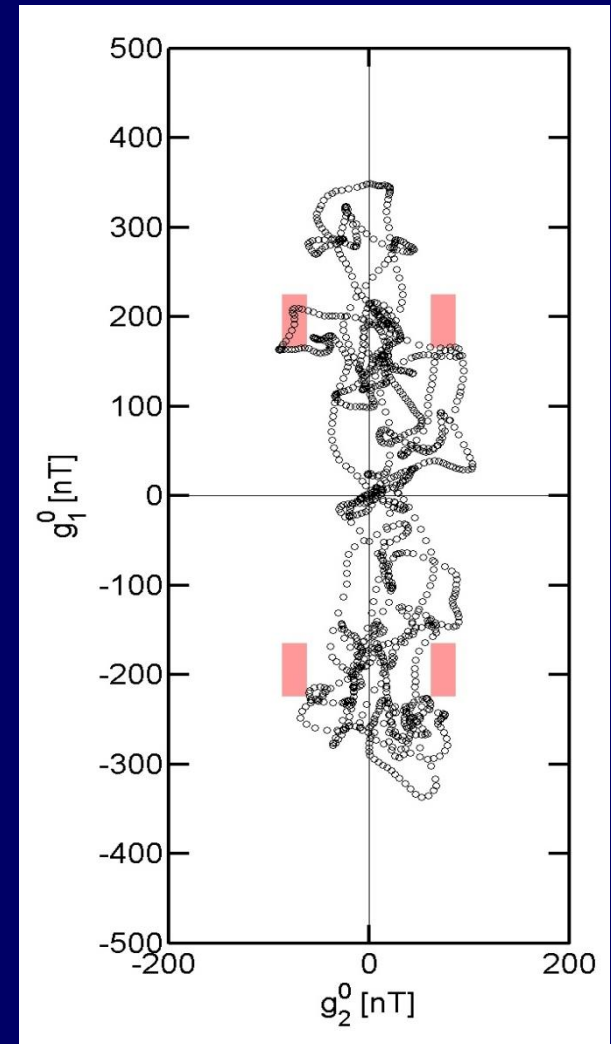
Christensen & Wicht, 2008

# Quadrupole / dipole ratio

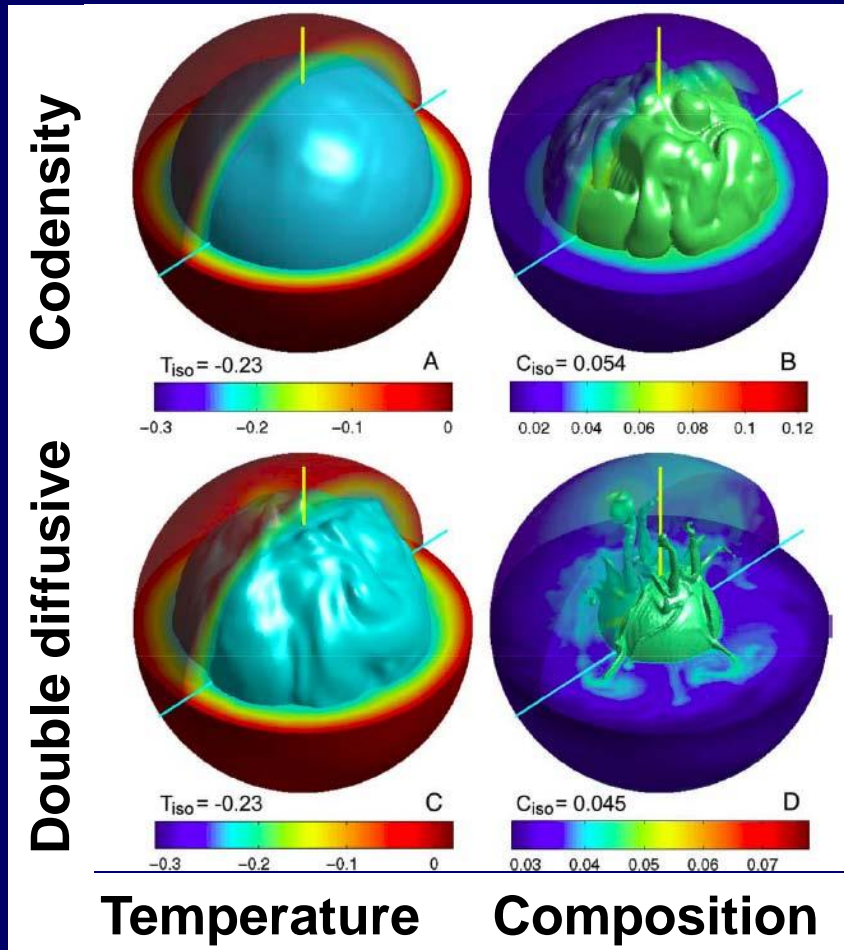


Snapshot fitting the present-day field of Mercury in terms of field strength and  $g_{20}/g_{10}$  ratio

One case re-analysed from Christensen & Wicht, 2008



# Mercury: double diffusive model



IC growth  $\Rightarrow$  light element flux  
Simple model: same diffusivity for heat and concentration: codensity

but

Compos diffusivity  $\ll$  Thermal diff  
 $\Rightarrow$  Double diffusive convection

With DDC, convective fingers penetrate into stable layer and cause magnetic induction.

Outside magnetic field too strong compared to observation.

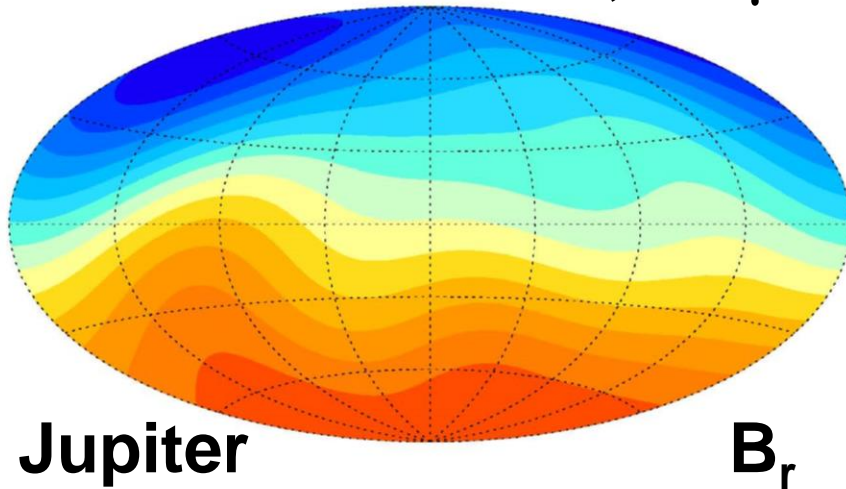
Compatible with observation when  $< 0.3\%$  sulphur in core

Manglik, Wicht & Christensen, 2010



# Saturn and Jupiter: Unlike siblings

$\pm 1,100 \mu\text{T}$



Jupiter's field 15xEarth's  
Similar proportion of  
non-zonal to zonal field  
as Earth. Dipole tilt 10°

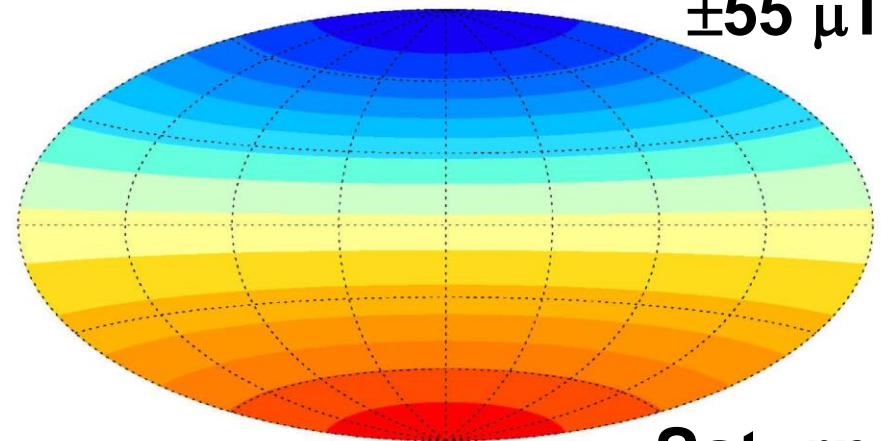
Saturn's field 0.7xEarth's

No non-zonal field  
detected so far.

Dipole tilt < 0.06°

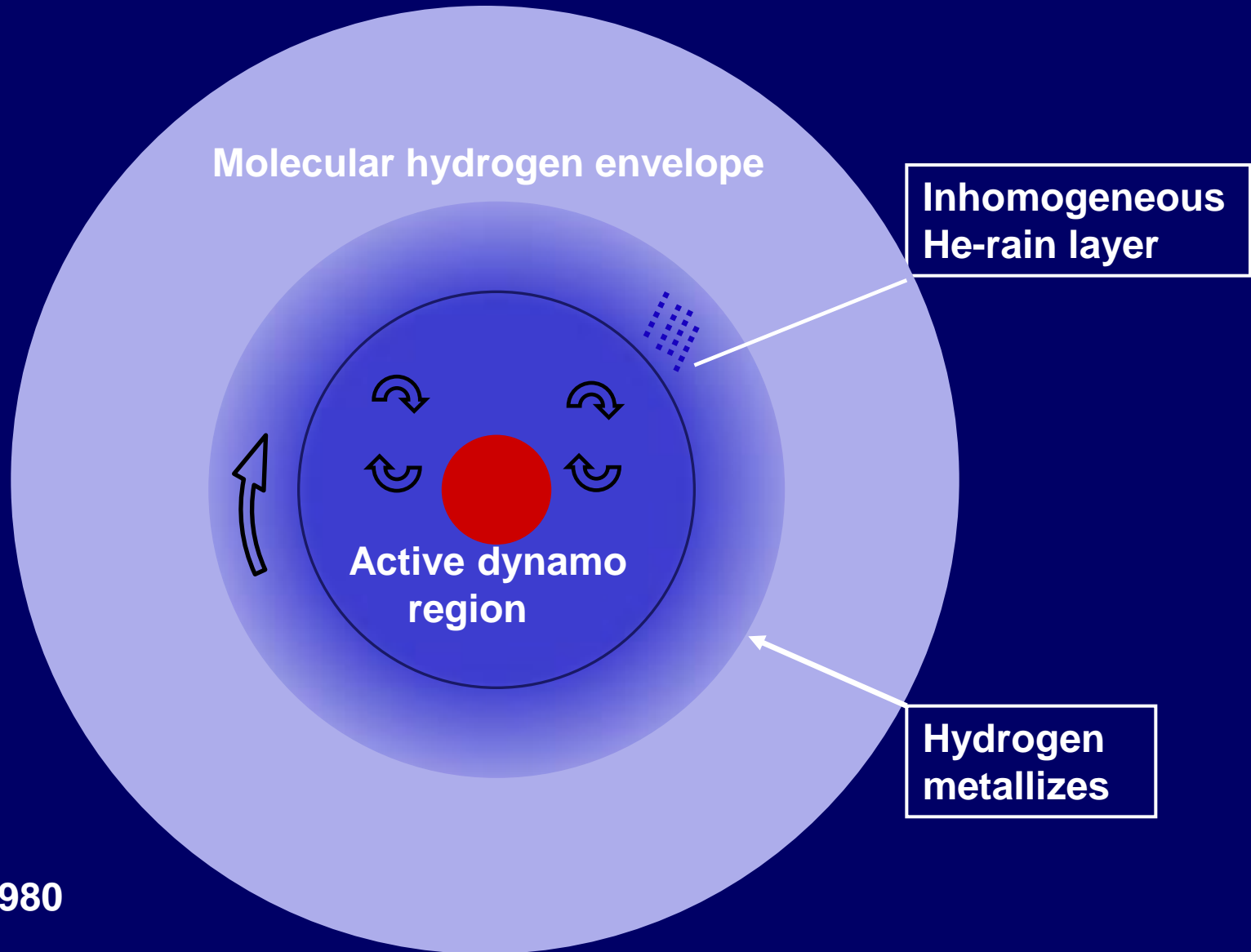
Cao et al., 2011

$\pm 55 \mu\text{T}$



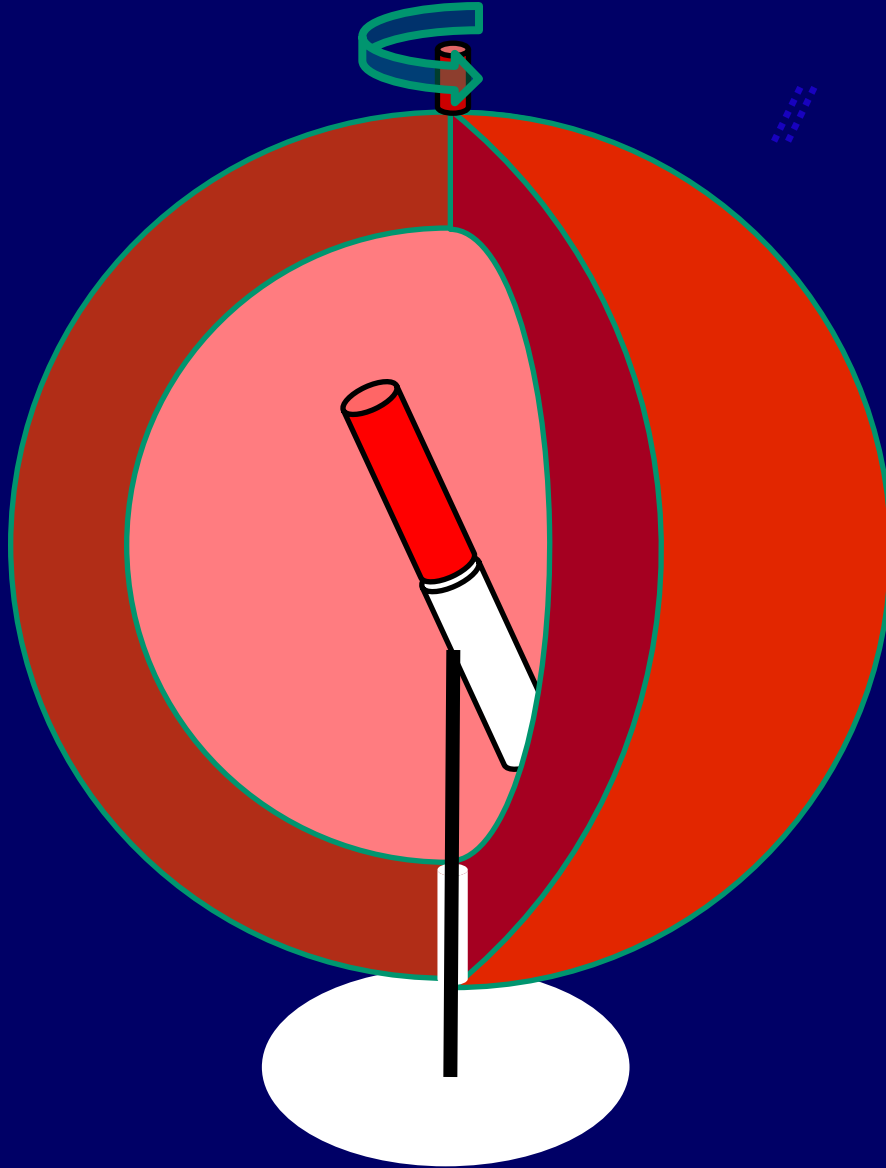


# Saturn: Dynamo below He-rain layer



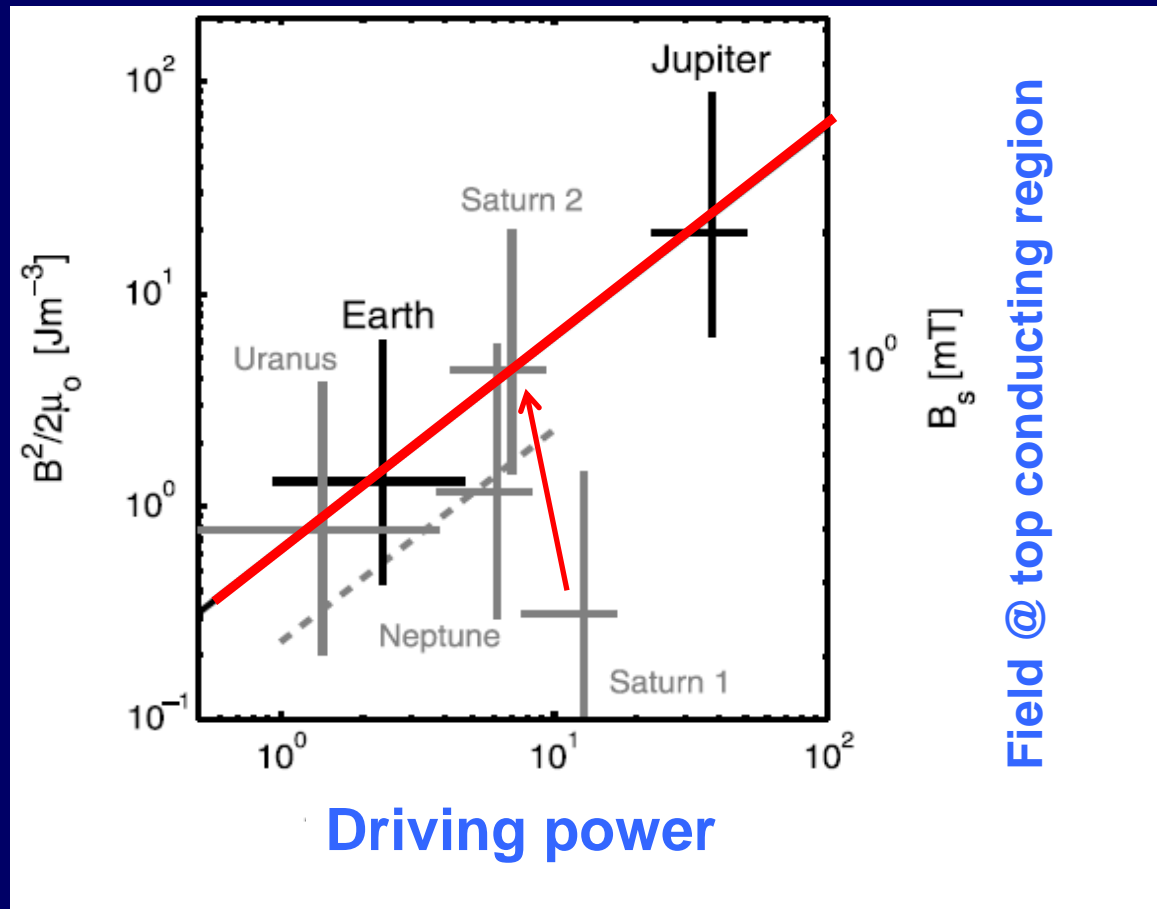
Stevenson, 1980

# A Gedankenexperiment



**Skin effect  
eliminates time-  
variable component  
(equatorial dipole)  
and lets static  
component (axial  
dipole) pass**

# Field strength for deep dynamo



Field @ top conducting region

With top of Saturn's dynamo at  $0.4 R_s$  rather than at  $0.62 R_s$  observed field complies with scaling relation Christensen, 2010

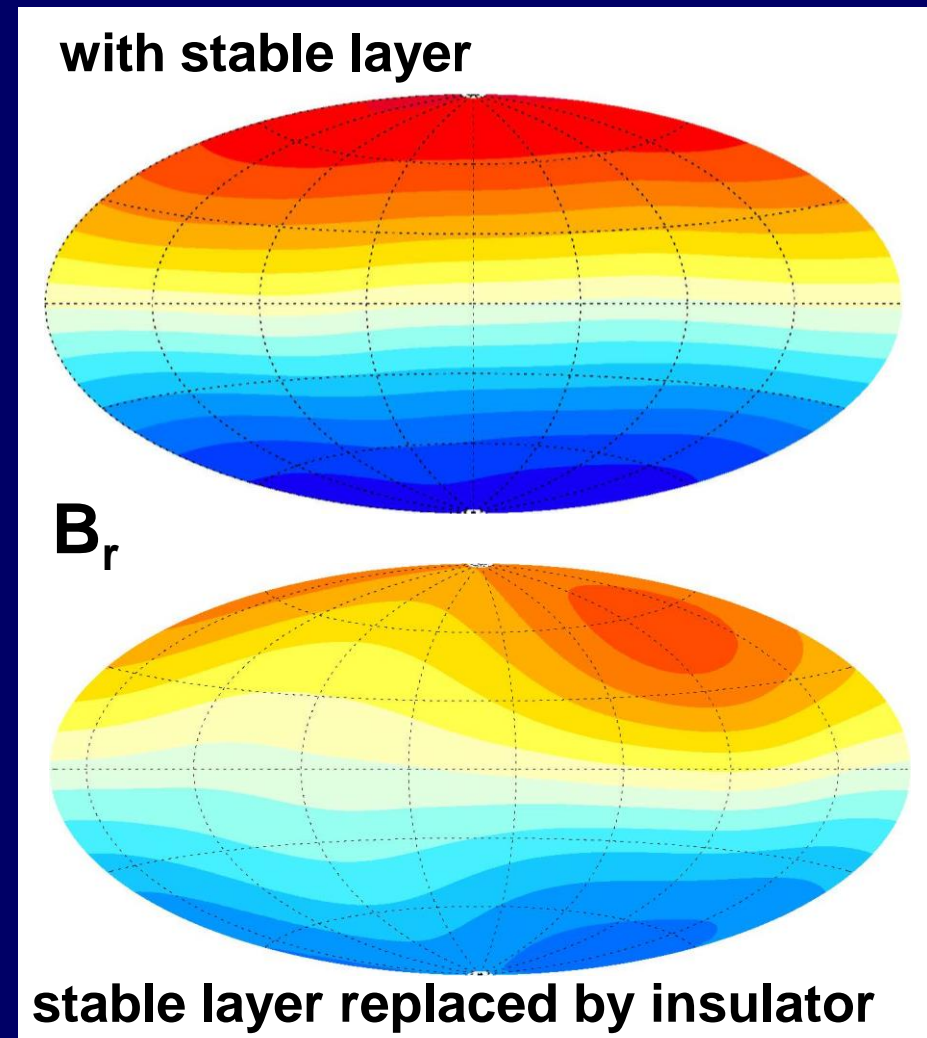
# Dipolar dynamo with stable layer

Strong non-zonal field component largely filtered out by stable layer.

Residual dipole tilt too large compared to observation ( $1.5^\circ$  vs.  $<0.06^\circ$ )

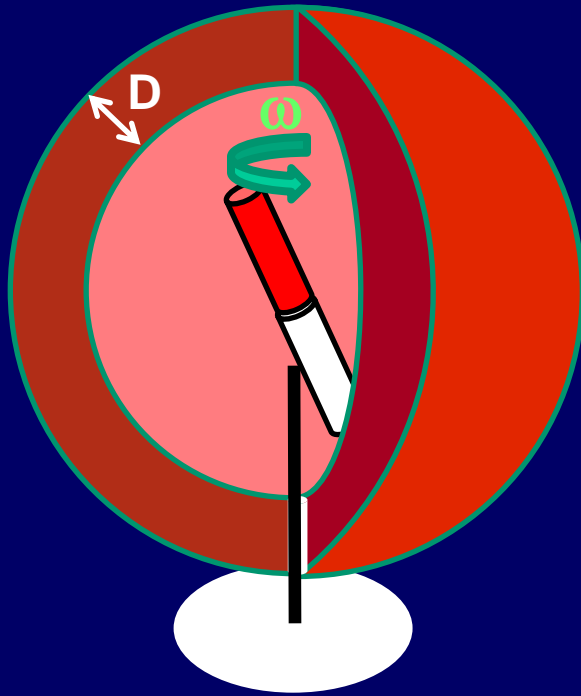
Model unable to match ratio of dynamo advection time to diffusion time @ stable layer

Christensen & Wicht, 2008





# Skin effect



Stagnant stable layer (thickness  $D$ )  
with magn. diffusivity  $\lambda$

Damping factor  $f_{\omega} = \exp(- [\omega D^2 / 2\lambda]^{1/2})$

With plausible values for  $\lambda$  and  $\omega$ , a  
thin layer ( $D=1000$  km, 1.5% of  $R_S$ )  
enough to damp by factor  $10^4$ .

Picture of equatorial dipole oscillating with single frequency  
too simple.

Consider power spectrum of dipole fluctuations.

# Skin effect

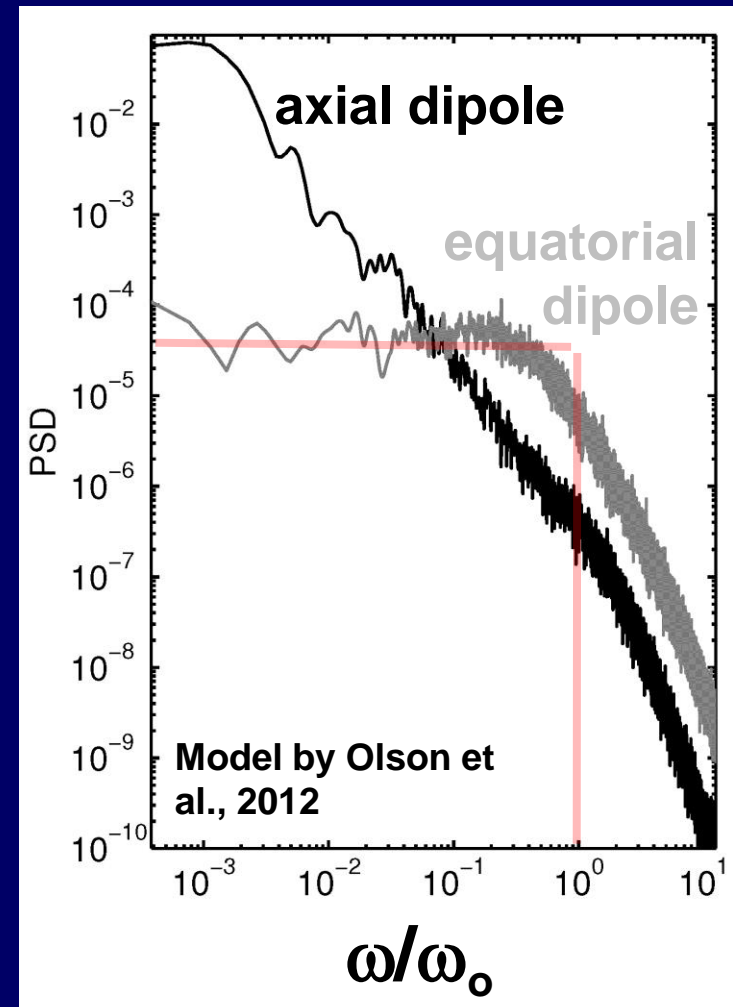
For equatorial dipole, temporal power spectrum flat up to  $\omega_o=U/R$  (inverse advection time in dynamo)

Folding white spectrum up to  $\omega_o$  with  $f_\omega$  and ignoring power at  $\omega>\omega_o$  results in net attenuation factor

$$f = (\lambda R / [2D^2U])^{1/2}$$

Saturn:  $U=1$  cm/s  $R=30,000$  km  
 $D=10,000$  km  $\lambda=2$  m<sup>2</sup>/s  $\Rightarrow f=1/170$

10° dipole tilt at top of dynamo  
reduced to 0.06° above stable layer

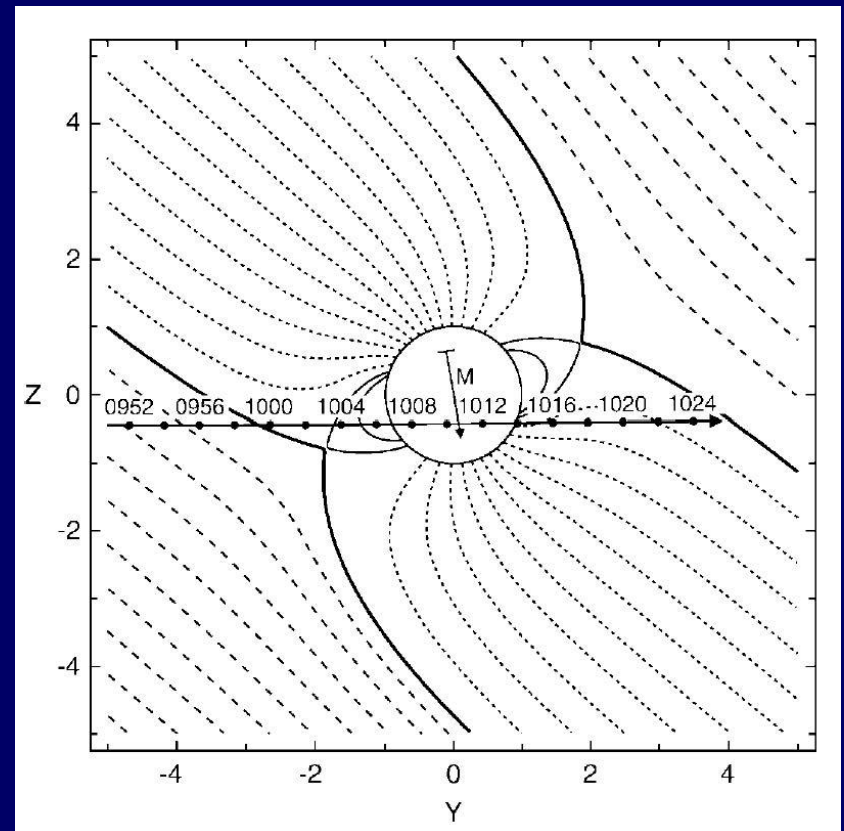


Christensen et al., 2015 (submitted)

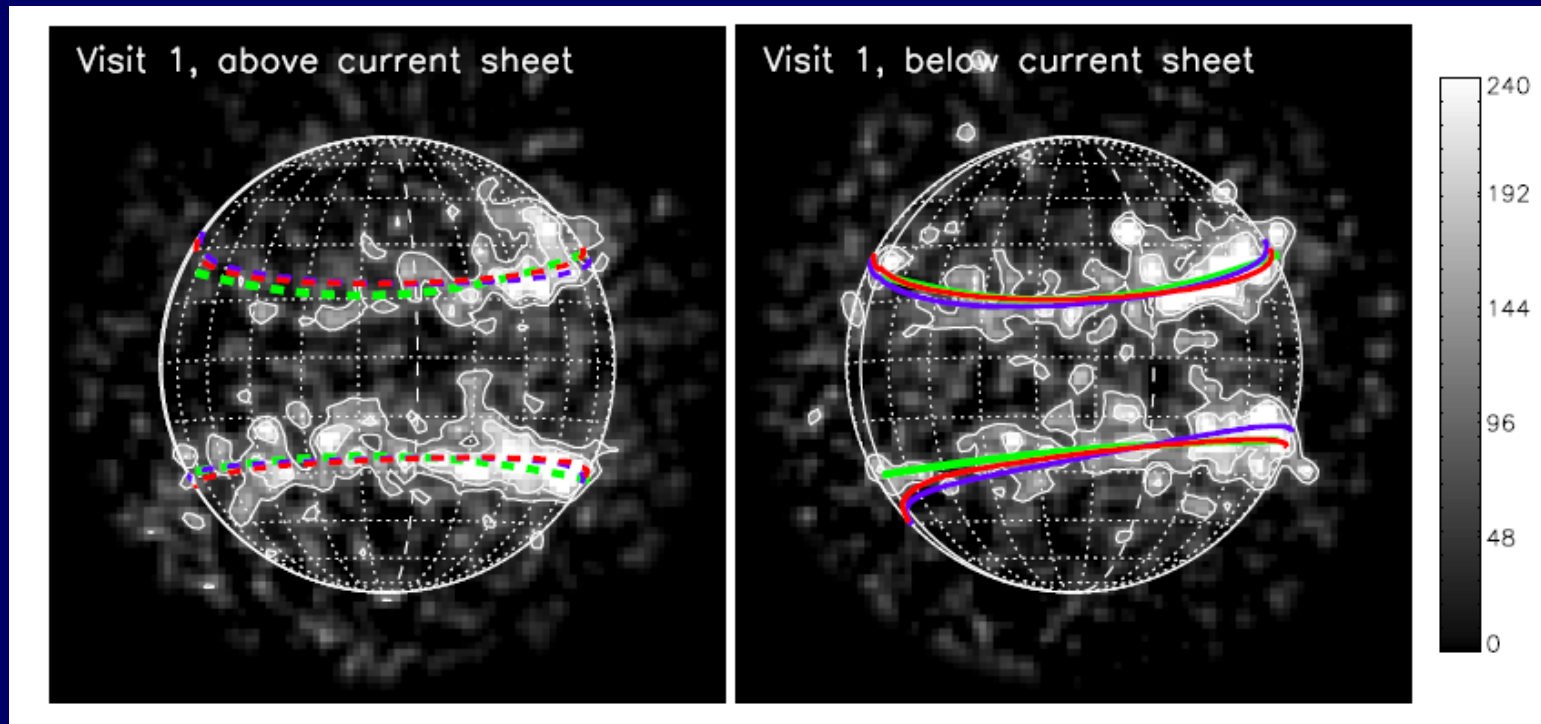
# Ganymede's magnetic field

- Galileo mission detected intrinsic magnetic field
- $g_{10}=711$  nT Dipole tilt  $4^\circ$
- Two field models equally consistent with data:
  - (1) Dipole + quadrupole
  - (2) Dipole + field induced in ocean, no quadrupole

Kivelson et al., 2002



# New evidence for induced field



**Aurora on Ganymede at open-closed fieldline boundary**  
**Rocking of aurora location with Jupiter's rotation in HST**  
**images less than expected for non-conducting planet**

Saur et al., 2015

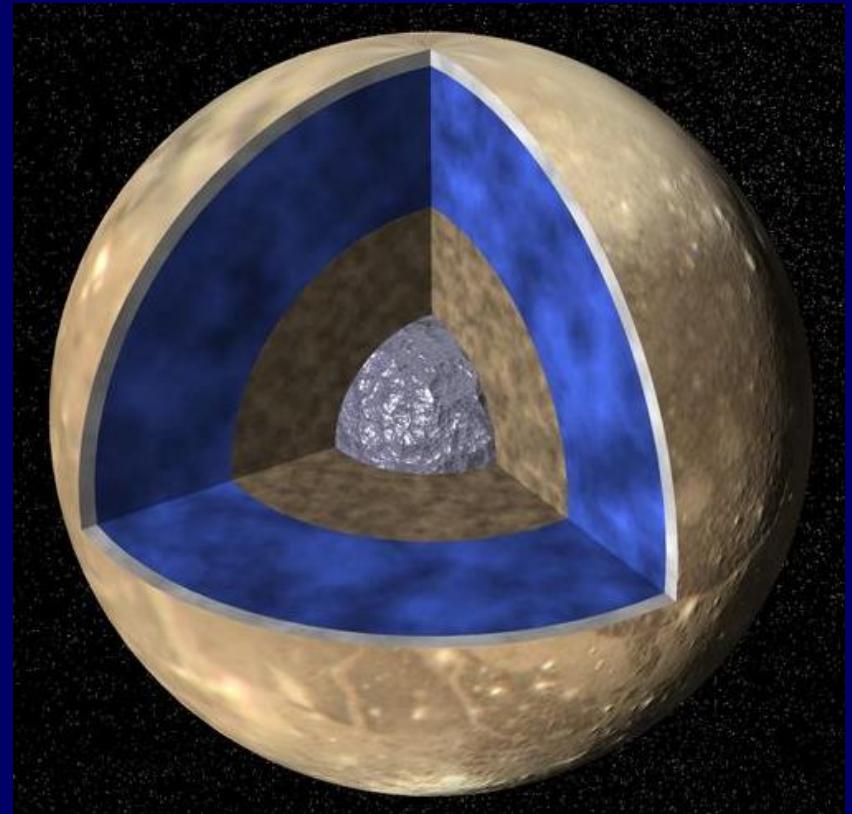


# Ganymede's dynamo field

- Internal field generated by dynamo in metallic core
- Take  $R_2/R_1$  of oceanless field model as upper limit
- Core radius  $r_c$  is  $1/4 - 1/3$  of planetary radius  $r_G$
- Even when calculated at  $r_c = 0.25 r_G$ , quadrupole is untypically small

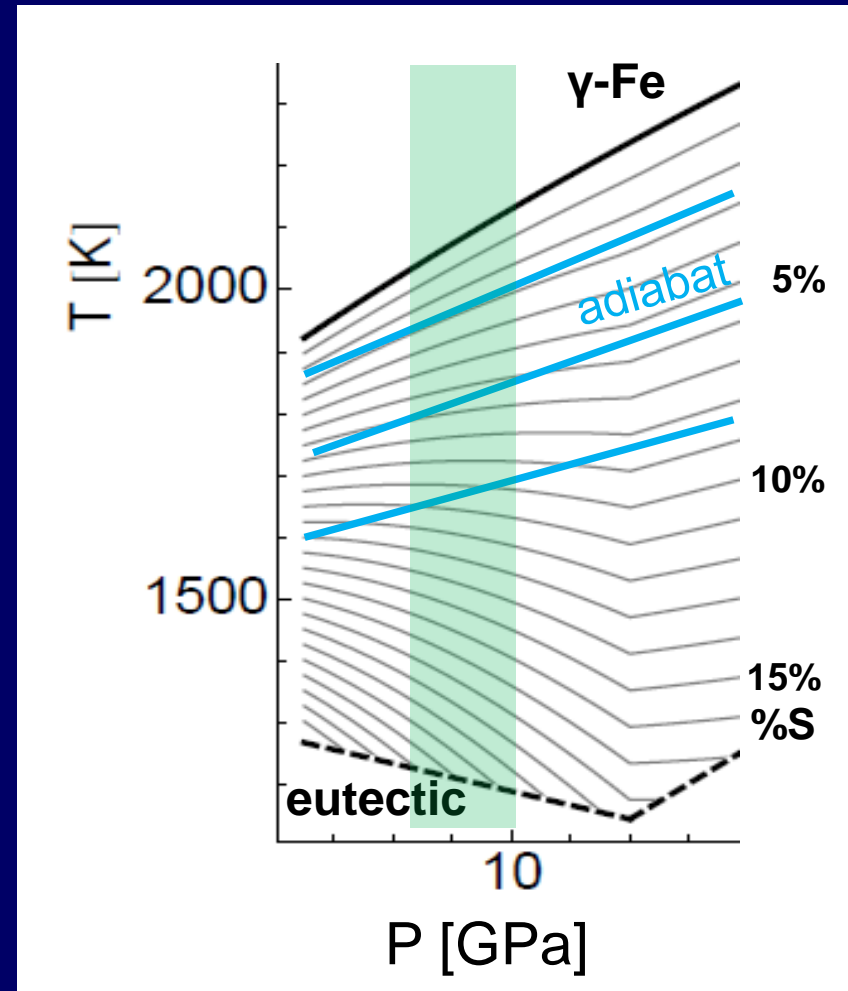
$$R_2 / R_1 < 0.04$$

(Earth CMB 2010:  $R_2/R_1 = 0.14$ )



# Fe – FeS melting curve

- Sulphur likely abundant in outer solar system
- S reduces melting temperature  $T_M$  and gradient  $dT_M/dP$
- In 5-10 GPa pressure range  $dT_M/dP < 0$  for  $>10\%$  S
- For more than few % S,  $dT_M/dP$  less steep than adiabatic T-gradient
- Crystallization of iron proceeds top-down

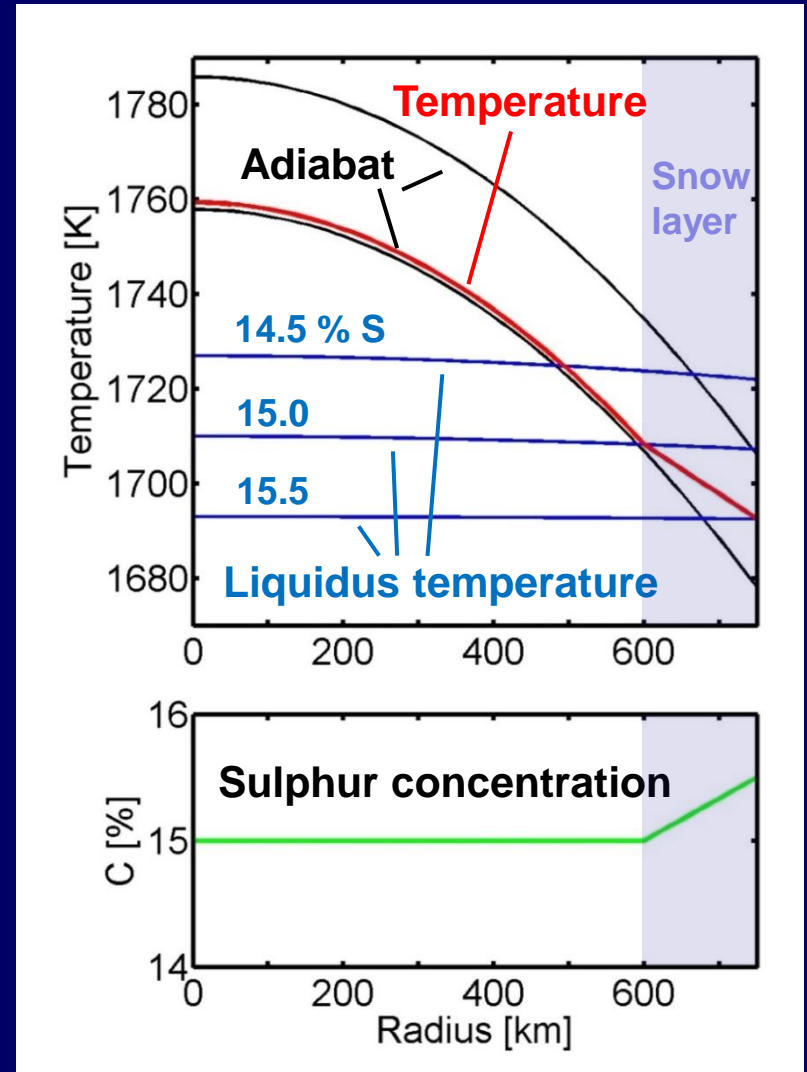


Hauck et al., 2006; Chen et al., 2008; Williams, 2009; Buono & Walker, 2010

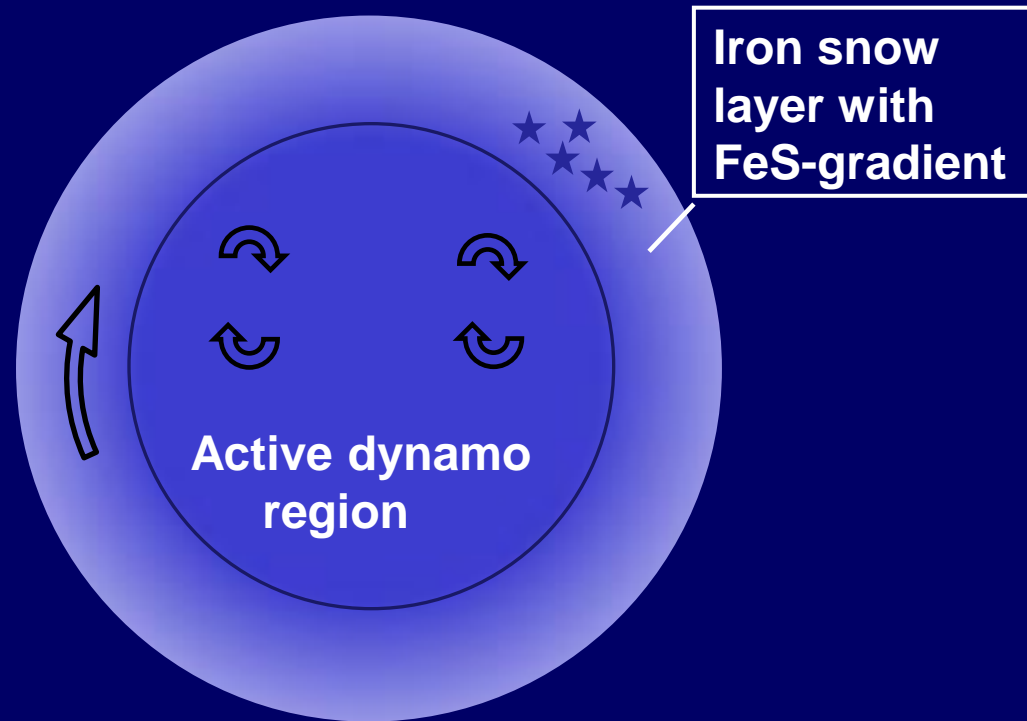
# Top-down crystallization

- **Adiabat steeper than melting point gradient**
- **Iron snow forms in top layer. Sinks and dissolves at bottom of this layer**
- **Sulphur-enrichment with stable compositional gradient in snow-forming layer. Here temperature everywhere at liquidus**

Hauck et al., 2006



# Dynamo below a snow layer

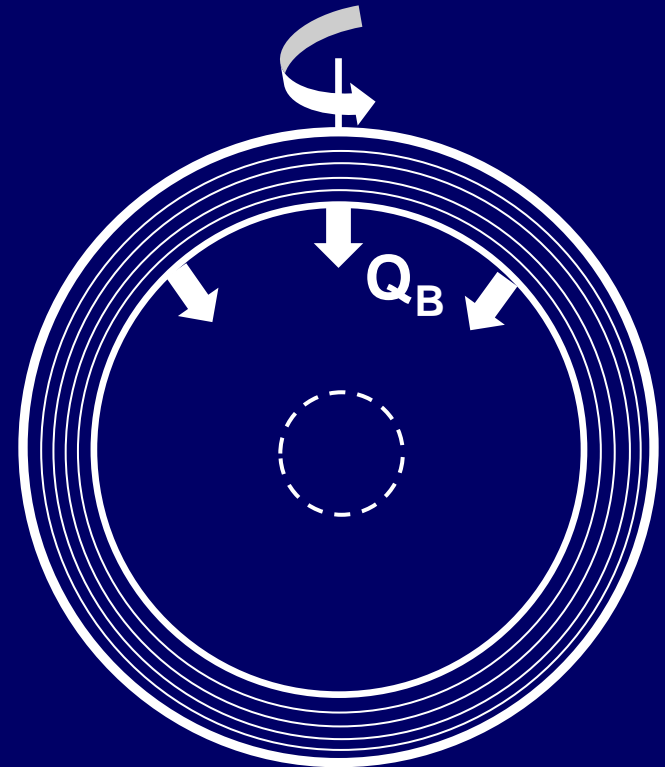


- **Fe-enrichment by melting of snow at top of interior region drives compositional convection**
- **Stable  $\Delta\rho$  in snow layer  $> 10^5$  times larger than typical  $\Delta\rho$  of convection  $\Rightarrow$  horizontal flow but no radial overturn**



# Ganymede dynamo models

- MHD dynamo model in rotating sphere, codensity, fixed buoyancy flux  $Q_B$  at top of dynamo region
- Reference model: **No extra outer layer**
- Snow layer model: **Strongly stable fluid layer above dynamo region. Only  $u_h$ , subject to viscous, inertial, Lorentz and Coriolis forces. Magnetic induction with  $u_r = 0$ ,  $u_h \neq 0$ .**

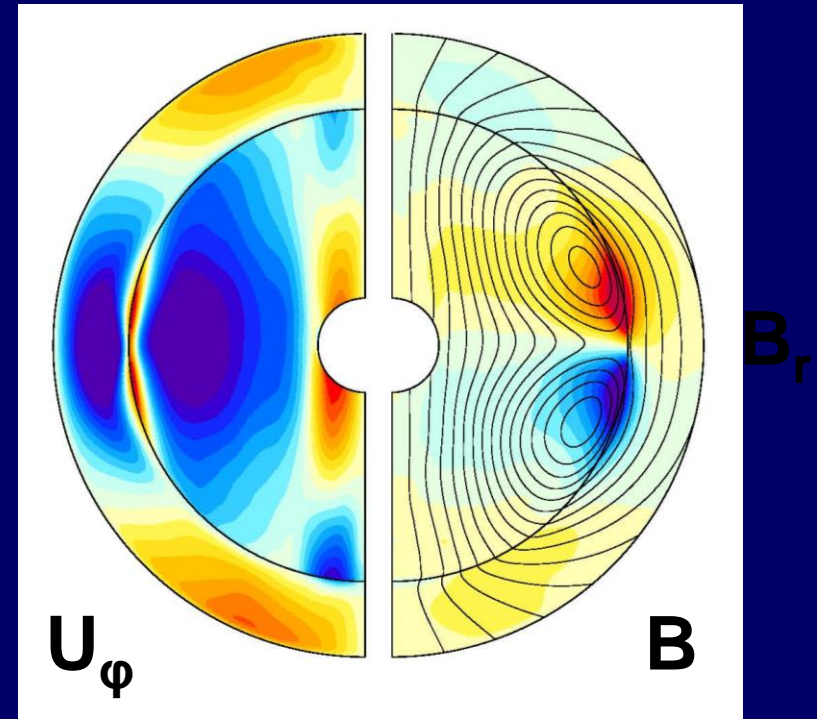
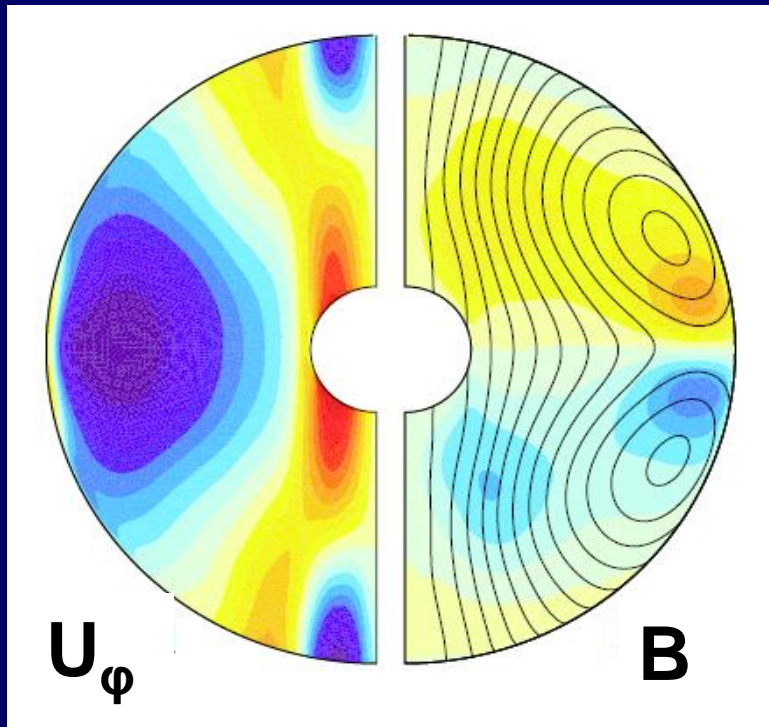


Christensen, 2015a

Christensen, 2015b (corrigendum)

# Azimuthal and time average

$$E = 3 \times 10^{-5} \quad Pm = 3 \quad Ra = 28 Ra_{crit}$$



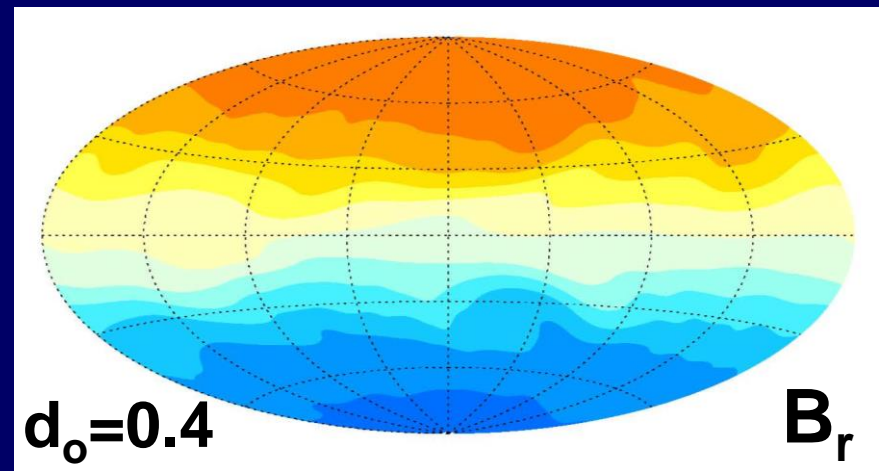
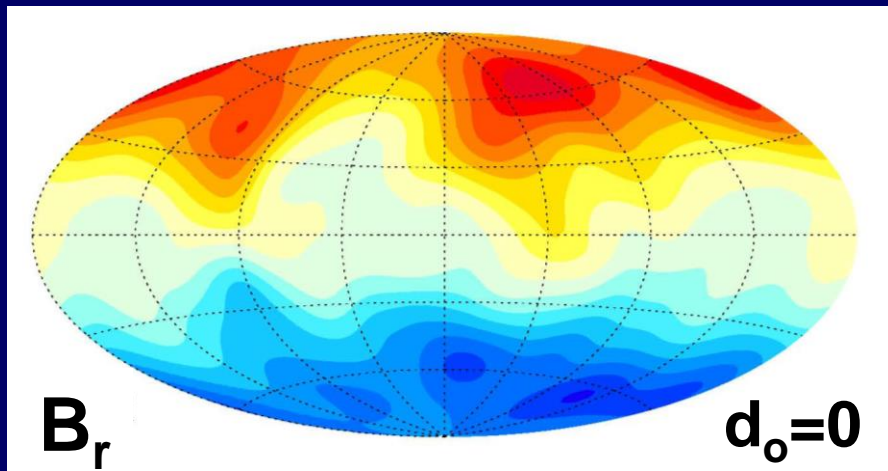
Reference model

Stable fluid layer

$$d_o = d / (R_c - d) = 0.4$$

# Field morphology

$$E = 3 \times 10^{-5} \quad Pm = 3 \quad Ra = 28 \times Ra_{crit}$$



**Reference model**

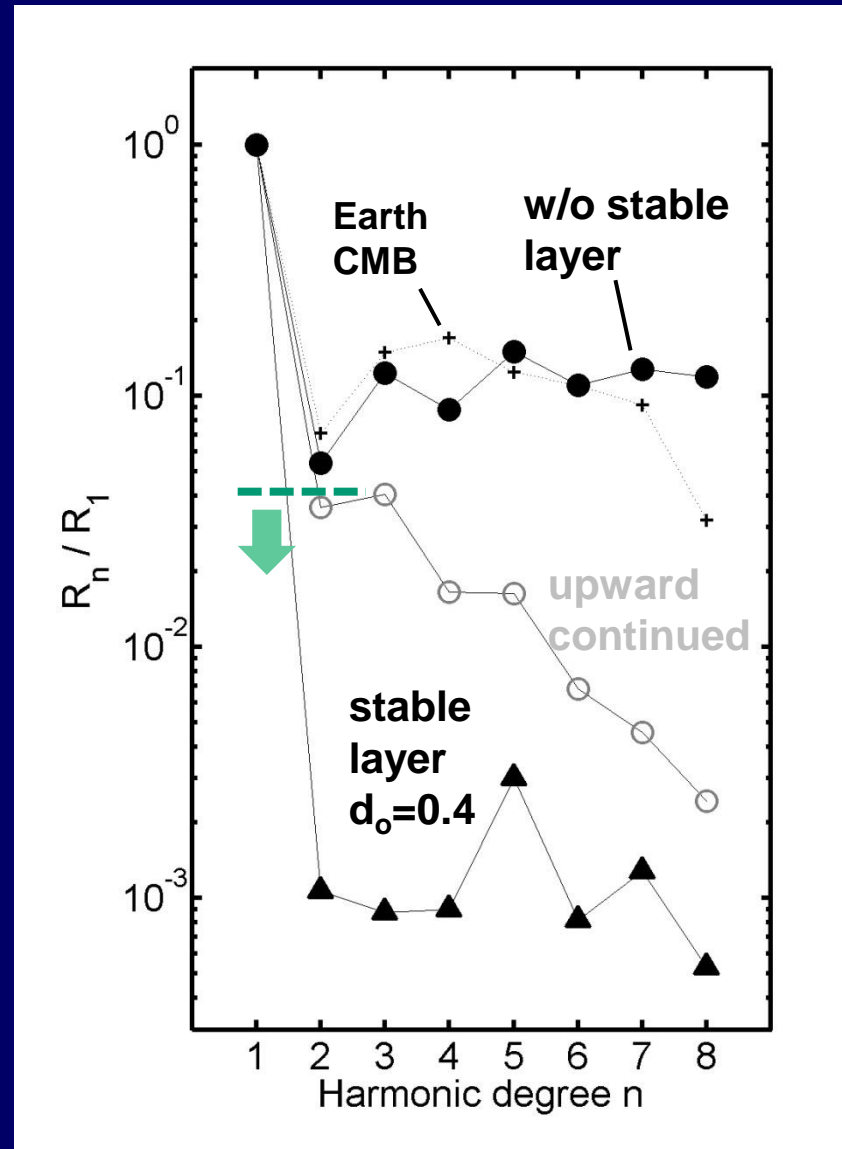
at  $r = r_c + 0.4$

**Stable fluid layer**

at  $r = r_c$

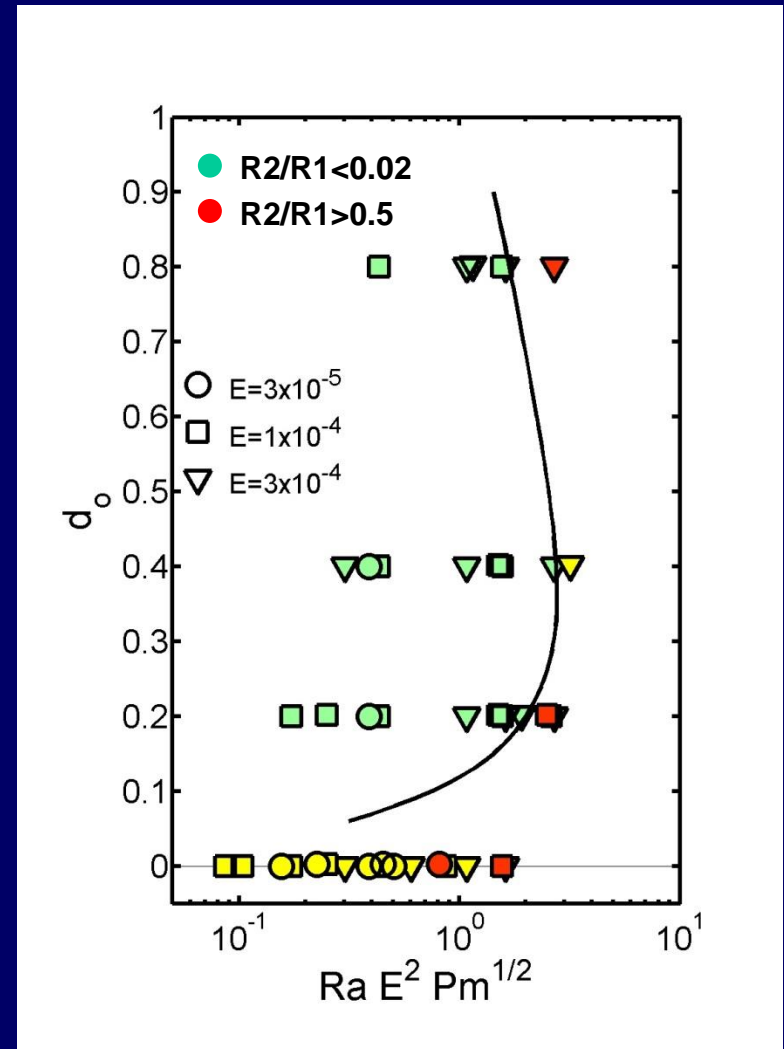
# Magnetic power spectra

- At top of the core, time-averaged, normalized with dipole
- $E=3 \times 10^{-5}$   $Pm=3$   $28Ra_c$
- Reference model w/o stable layer similar to geodynamo spectrum
- With snow layer drop in multipoles
- $R_2 / R_1 \sim 0.001$



# Dynamo regime

- Empirical ordering parameter  $Ra E^2 Pm^{1/2}$
- Dipolar dynamos with low  $R_2/R_1$  for stable layer thickness  $d_o \geq 0.2$  ( $\approx 100$  km) and  $Ra E^2 Pm^{1/2} < 2$
- Ganymede value:  
 $Ra E^2 Pm^{1/2} \sim O(1)$



# Field strength: How to rescale?

$$B = B^* \sqrt{\rho \mu \lambda \Omega} \quad t = t^* D^2 / \nu$$

$$Ra = \frac{g Q_B D^4}{\rho \nu \kappa^2} \quad E = \frac{\nu}{\Omega D^2} \quad Pm = \frac{\nu}{\lambda} \quad Pr = \frac{\nu}{\kappa}$$

- Take true values of  $D, \rho, g$
- To match non-dimensional model parameters, set diffusivities  $\nu, \kappa, \lambda$  to whatever is required
- We must also compromise either on the buoyancy flux  $Q_B$  (overdrive) or on  $\Omega$  (underrotate)
- Scaling studies suggest that  $B$  depends on  $Q_B$ , but not strongly on diffusivities or  $\Omega$  (Christensen & Aubert, 2006)
- **Conclusion: take correct  $Q_B$ , compromise on  $\Omega$**



# Field strength

$$E = 3 \times 10^{-5}$$

$$Pm = 3$$

$$Pr = 1$$

$$Ra = 2.5 \times 10^8$$

Fix  $D$ ,  $\rho$ ,  $g$  to Ganymede,  $Q_B = 5000 \text{ kg/s}$

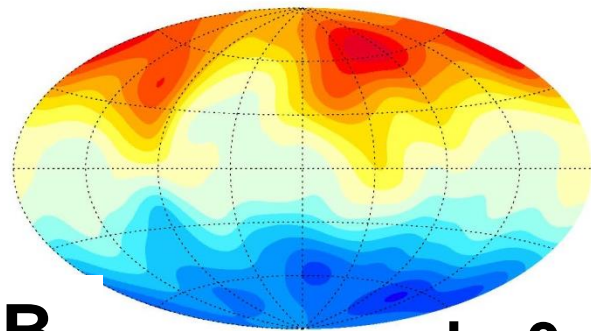
Rückriemen  
et al., 2015

➤  $\nu = \kappa = 3.8 \text{ m}^2/\text{s}$

$$\lambda = 1.3 \text{ m}^2/\text{s}$$

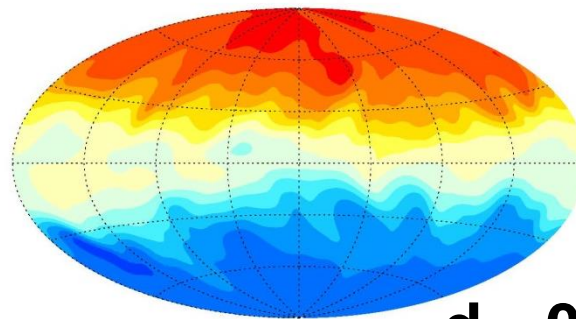
$$2\pi/\Omega = 160 \text{ d}$$

$$g_{10} = 1030 \text{ nT}$$



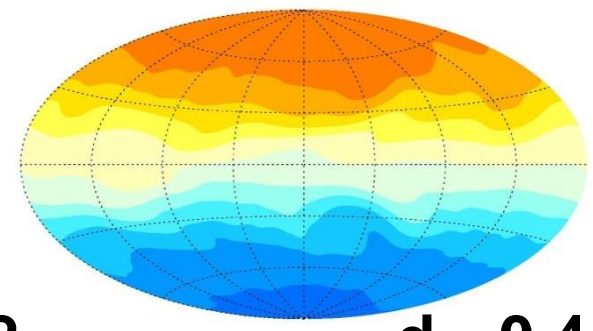
$$d_o = 0$$

$$g_{10} = 620 \text{ nT}$$



$$d_o = 0.2$$

$$g_{10} = 440 \text{ nT}$$



$$d_o = 0.4$$

(Observed  $g_{10} = 711 \text{ nT}$ )

# Conclusions

- **Stably stratified layer above dynamo region in planets: Mercury, Saturn, Ganymede (Earth?)**
- **Stable layer tends to make field more axisymmetric (small dipole tilt), lowers dipole moment**
- **Reduced quadrupole if dynamo field dipole-dominated. Multipolar dynamo field permits large axial quadrupole**
- **Test of concept by future observations: Skin effect prolongs time-scale of secular variation**