Medium-range Forecast Error: Air-sea Coupling, Clouds, and Scale of Meridional Propagation

Arindam Chakraborty

Centre for Atmospheric and Oceanic Sciences
Indian Institute of Science
Bangalore 560012, INDIA.

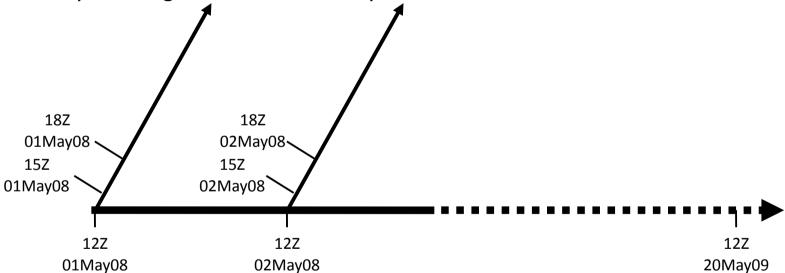
Background

- An accurate forecast of precipitation over the South Asian monsoon region at high spatial resolution up to 10 days in advance and beyond is necessary for decision making in different areas like agriculture and water resources management.
- □ Although significant improvements have been made in short-range forecasting by numerical models during the past few decades (Simmons and Hollingsworth 2002; Harper et al. 2007; Kalnay et al. 1998), there still exists large errors in short-to-medium range forecasts.
- ✓ Since May 2008, the WCRP and the World Weather Research Programme/ The Observing System Research and Predictability Experiment (WWRP/ THORPEX) have been organizing coordinated observations, modeling and forecasting organized tropical convection with a goal to better understand its multiscale structure and interaction.
- ➤ This program is titled as the Year of Tropical Convection (YOTC).
- As a part of this program, the European Centre for Medium-Range Weather Forecasts (ECMWF) generated high-resolution model forecasts up to 240 h from May 2008 through April 2010.

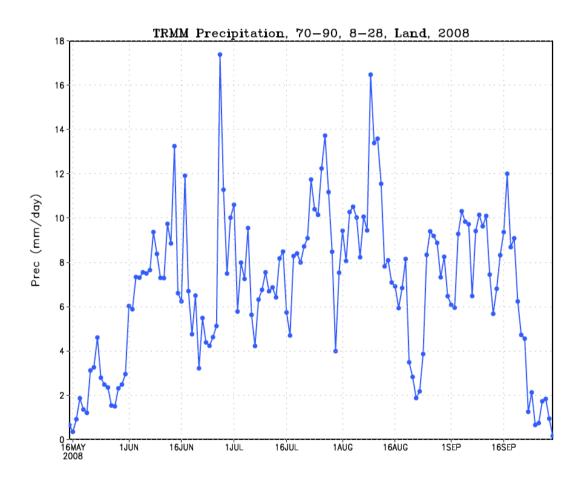
Model Forecast Details

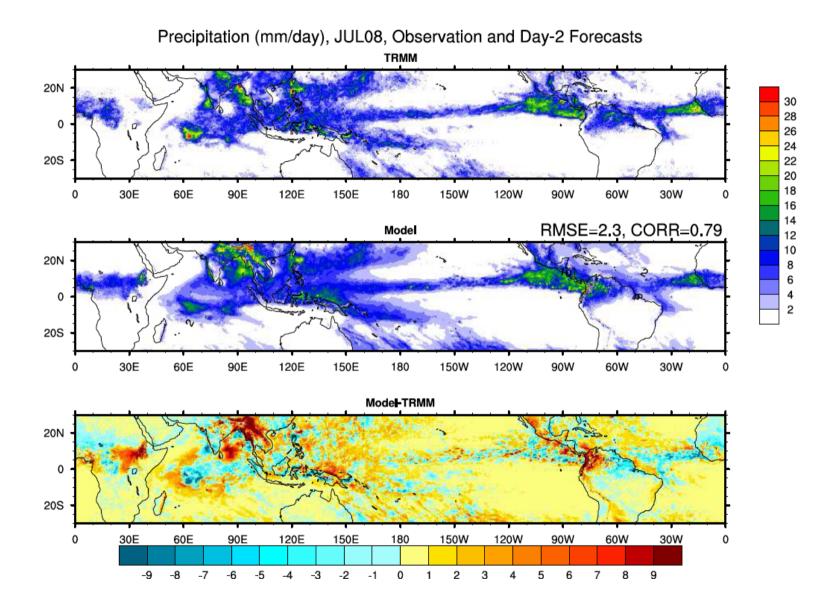
- > 3-hourly forecasts up to 96 hours, 6-hourly forecasts up to 240 hours.
- Forecasts started at 12 GMT of each day during 1 May 2008 through 20 May 2009.
- ➤ Average of 3 to 12 hour forecasts is termed as day-1 (calendar day of the forecast initial condition).
- > Average of 15 to 36 hour forecasts is termed as day-2.

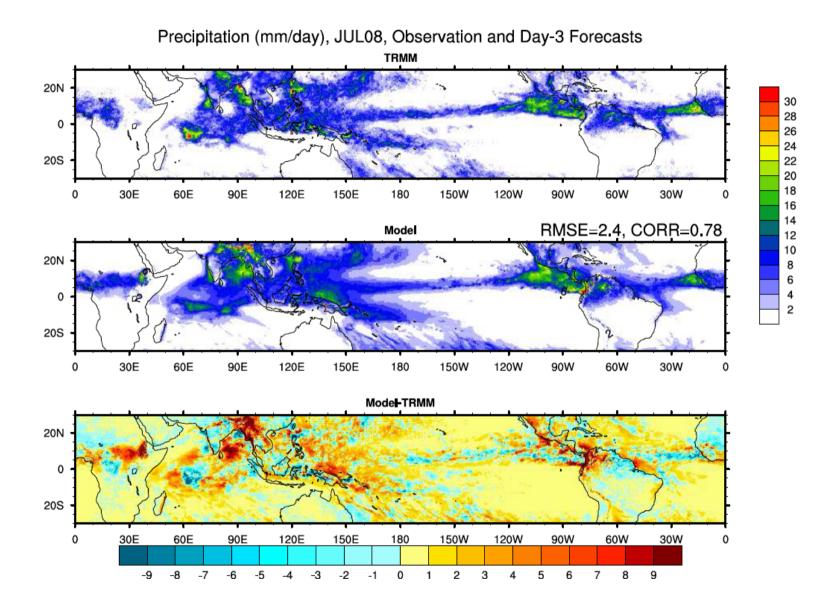


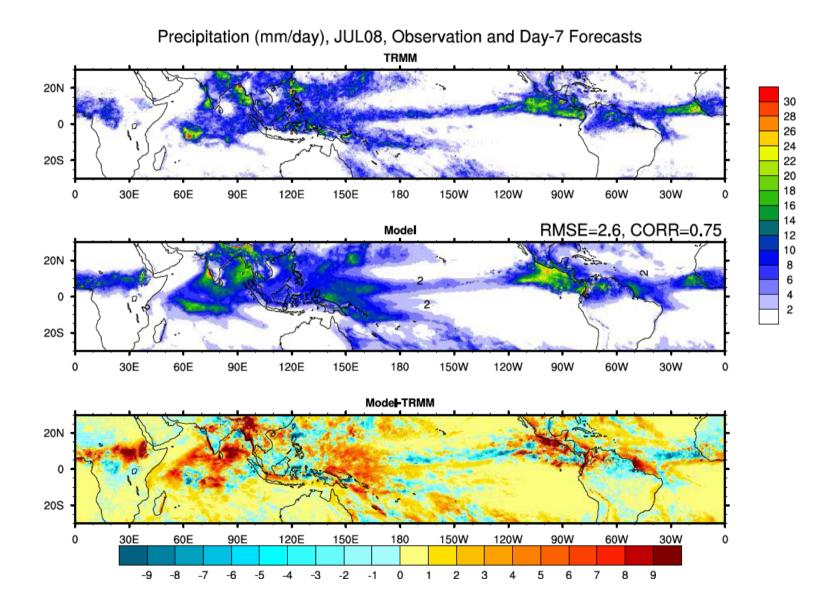


Observed Daily Precipitation over the Indian Region from TRMM Satellite Estimation

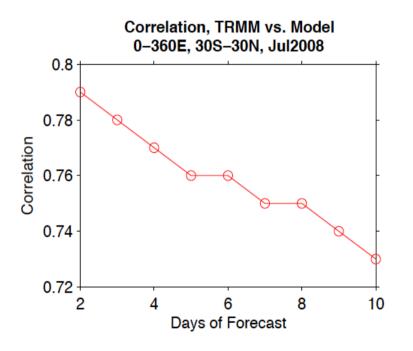


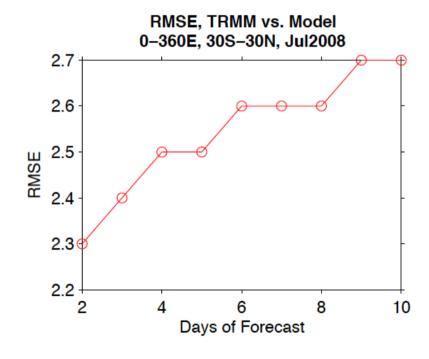


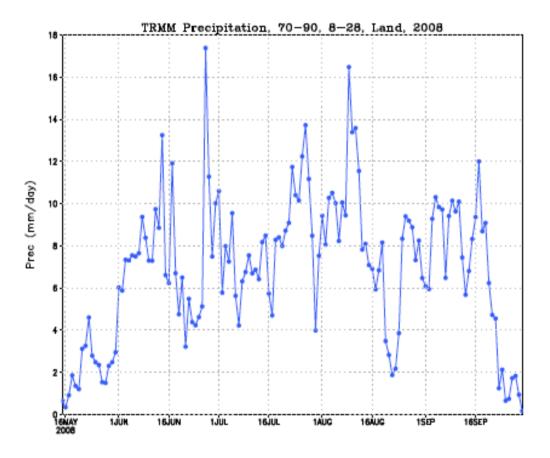


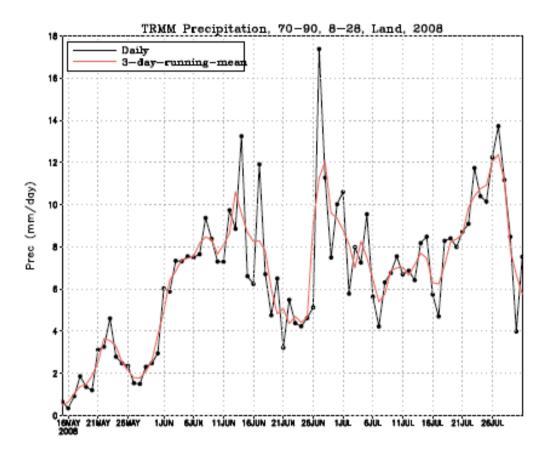


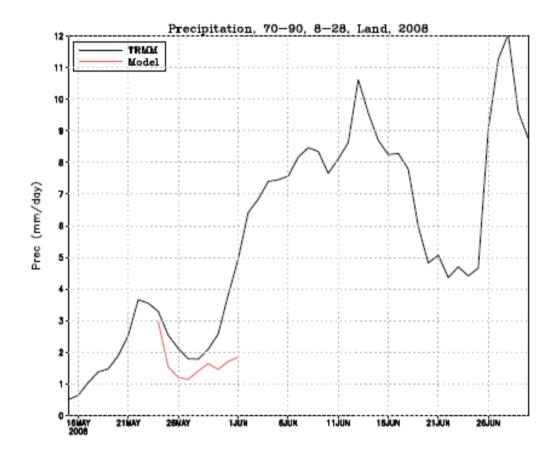
Skill Scores during July over the Tropics

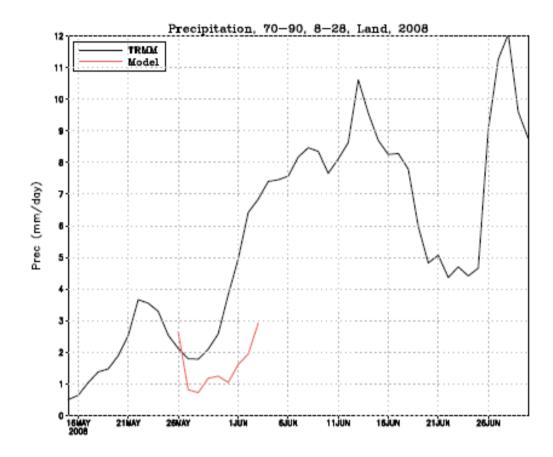


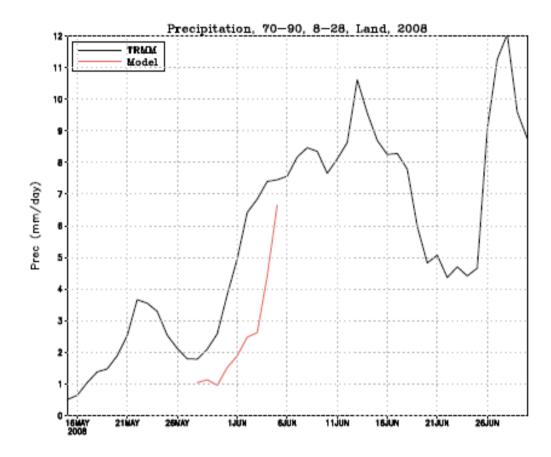


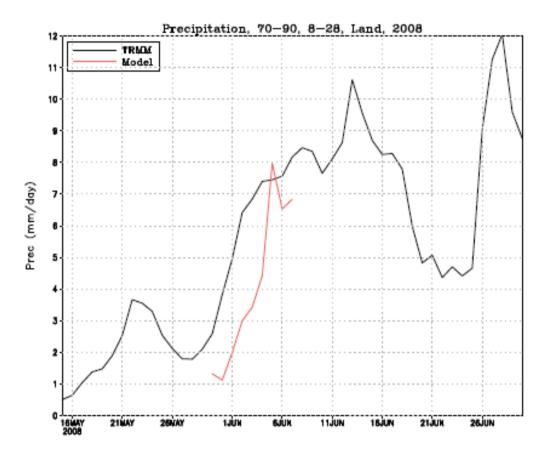


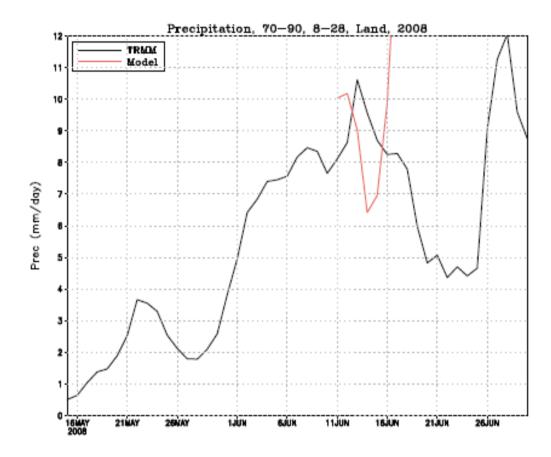


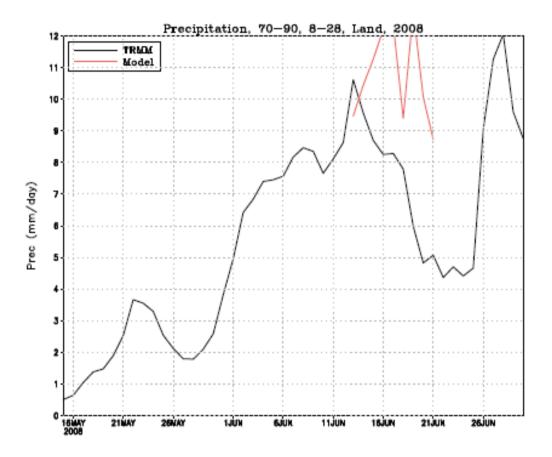


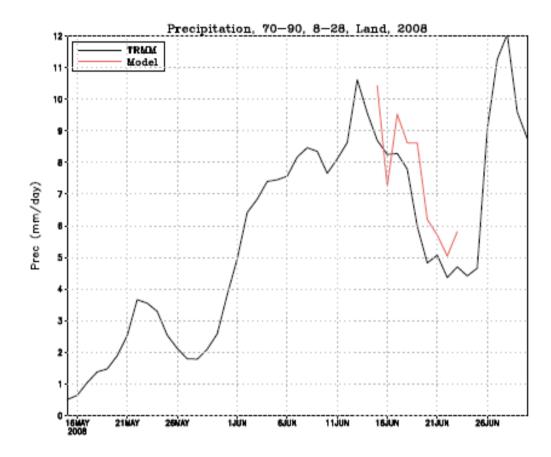


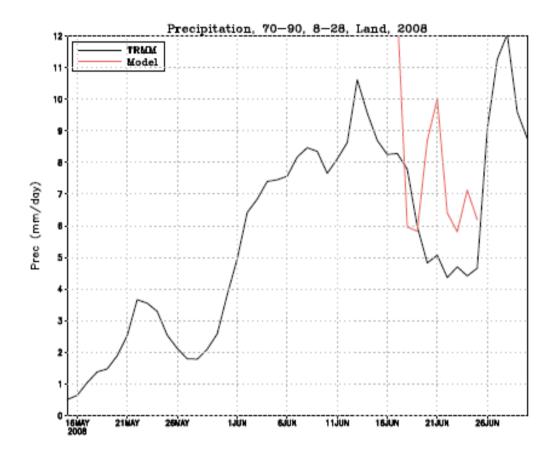


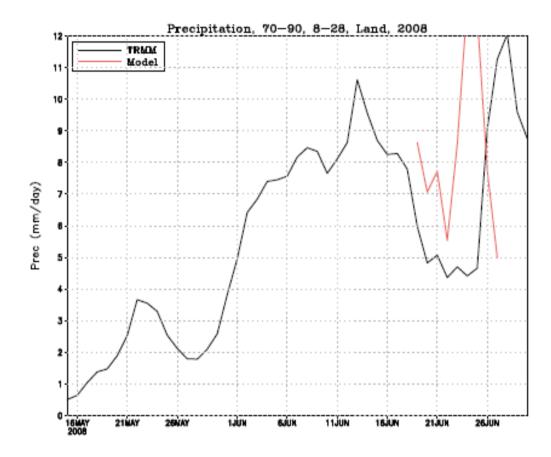


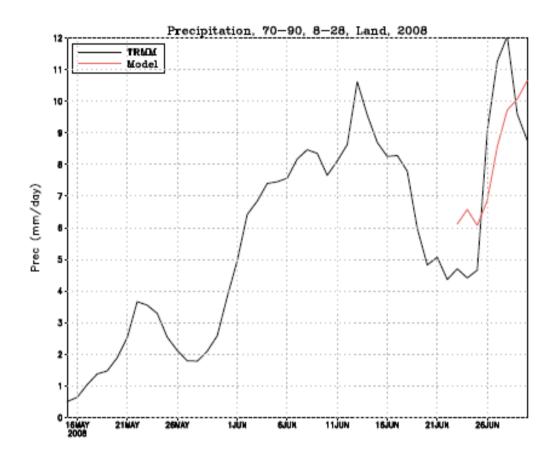


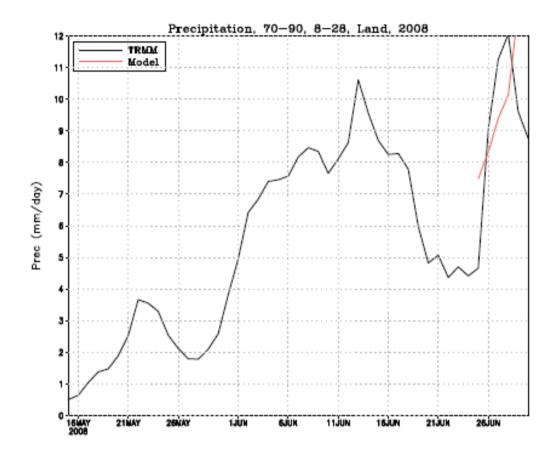


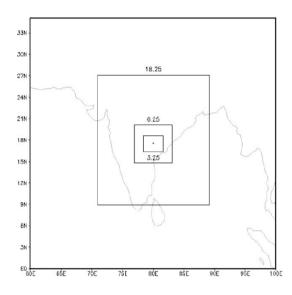


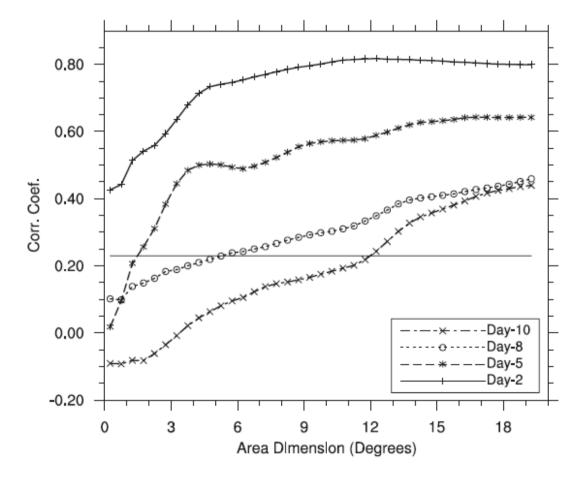


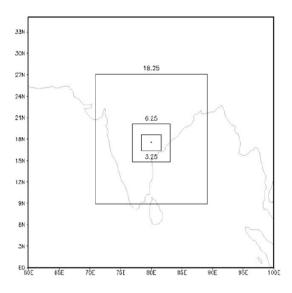


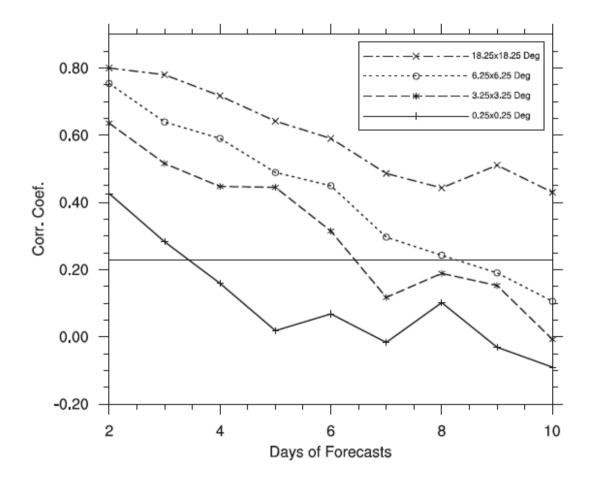


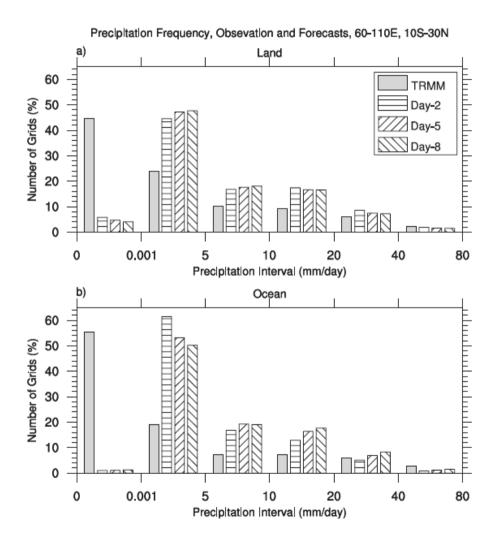




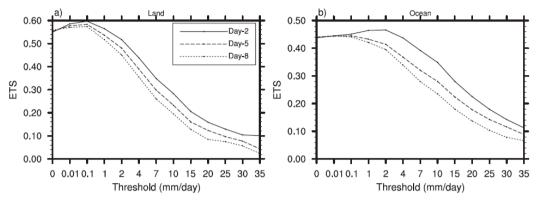








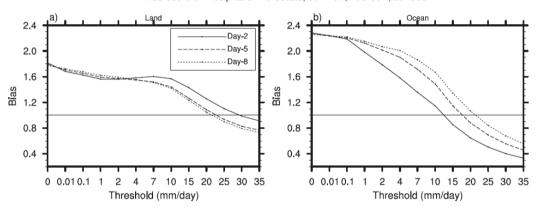
Equitable Threat Score of Precipitation Forecasts, 60-110E, 10S-30N, JJAS08



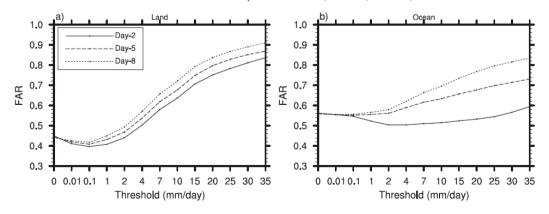
$$ETS = \frac{H - E}{F + O - H - E}$$

Bias Score of Precipitation Forecasts, 60-110E, 10S-30N, JJAS08

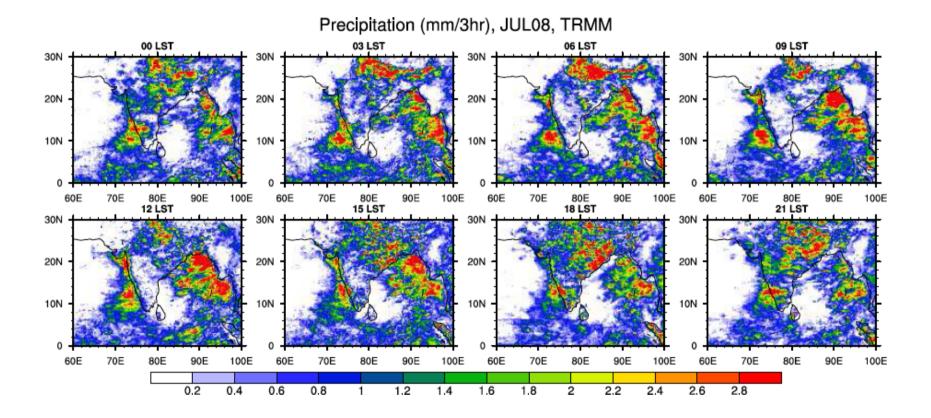
$$BIAS = \frac{F}{O}$$

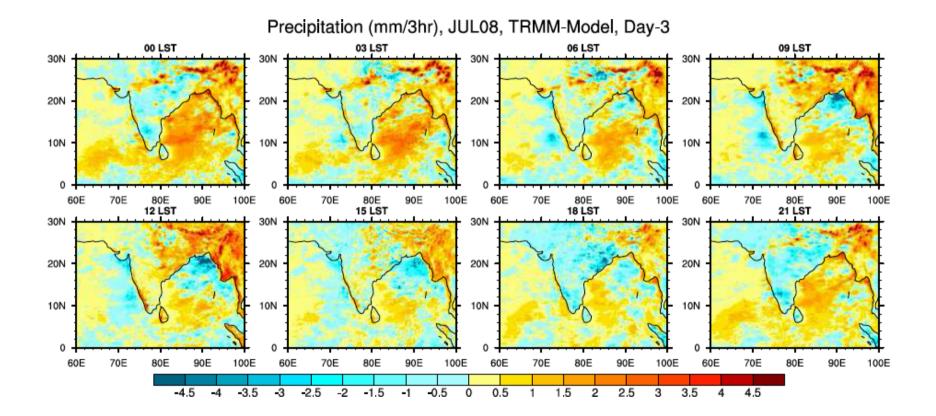


False Alarm Ratio of Precipitation Forecasts, 60-110E, 10S-30N, JJAS08

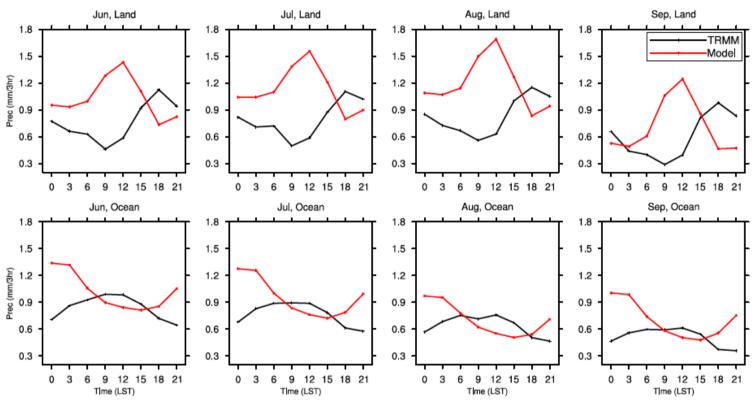


$$FAR = \frac{F - H}{F}$$

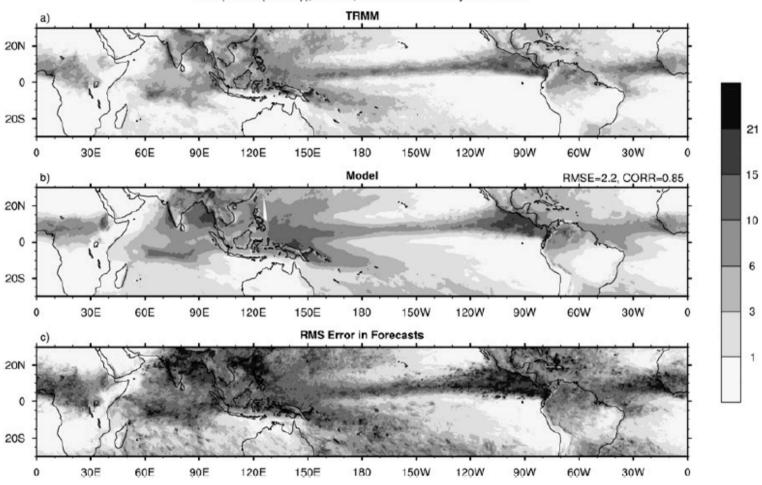




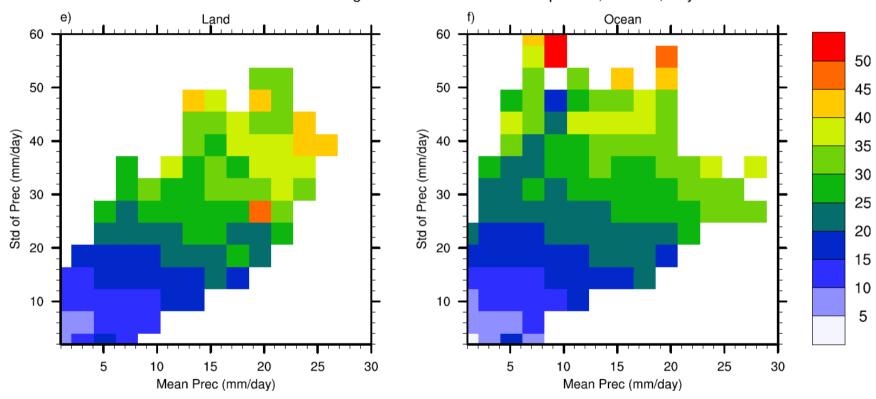
Precipitation (mm/3hr), TRMM and Day-3 Forecasts of Model



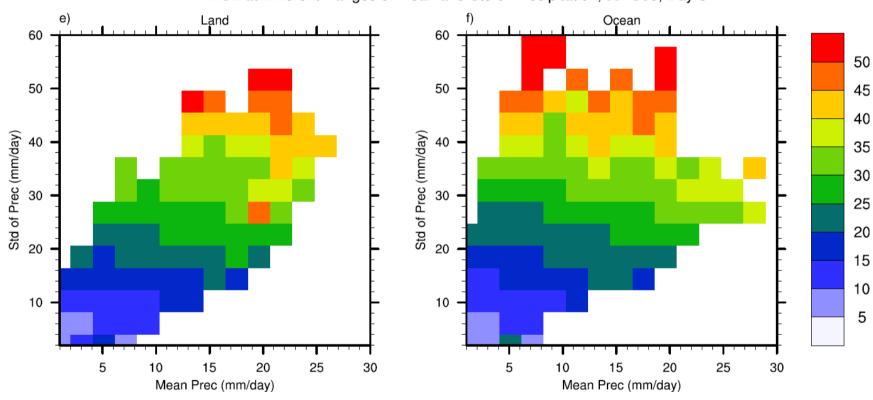
Precipitation (mm/day), JJAS08, Observation and Day-5 Forecasts



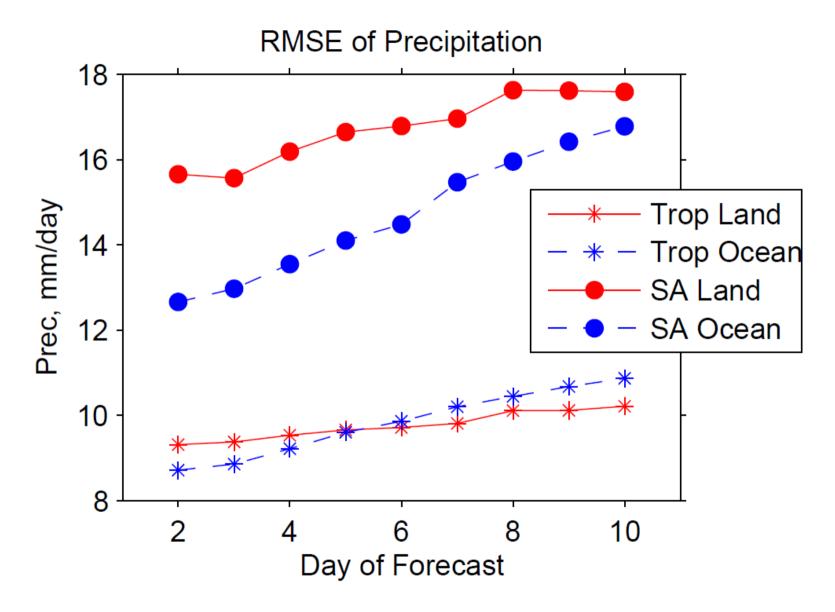
RMSE at Different Ranges of Mean and Std of Precipitation, JJAS08, Day-2

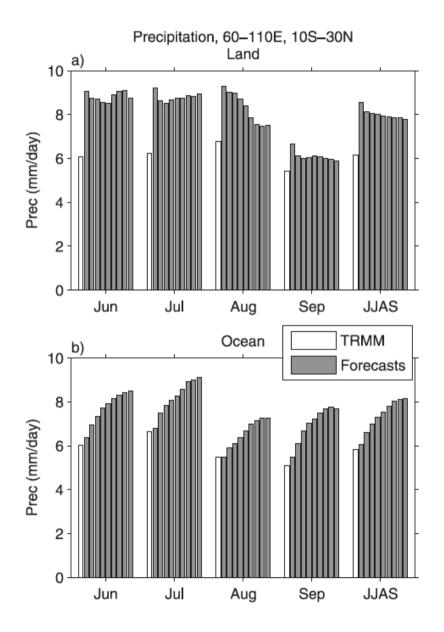


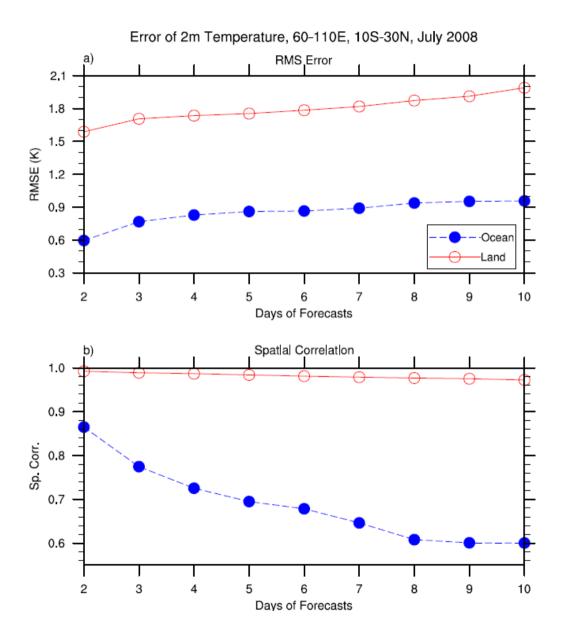
RMSE at Different Ranges of Mean and Std of Precipitation, JJAS08, Day-5



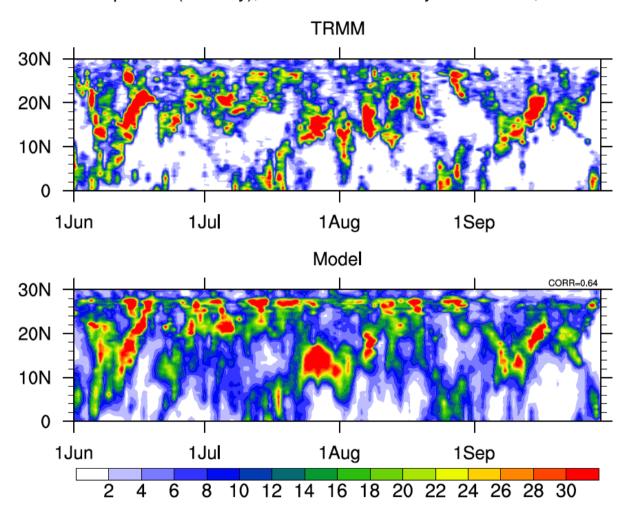
RMSE at Different Ranges of Mean and Std of Precipitation, JJAS08, Day-8 Land Std of Prec (mm/day) Std of Prec (mm/day) Mean Prec (mm/day) Mean Prec (mm/day)



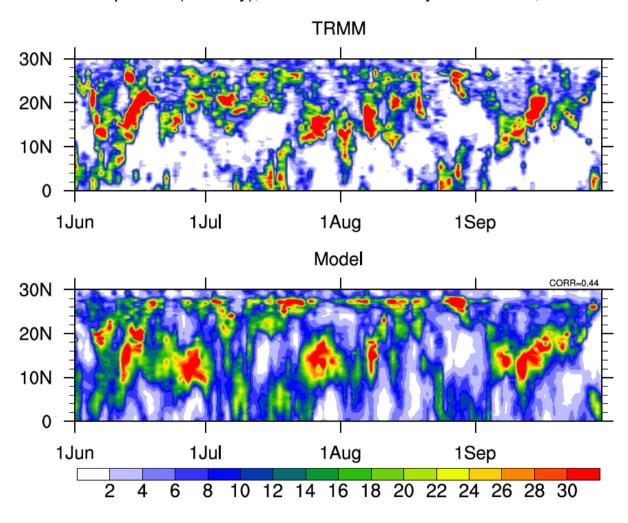




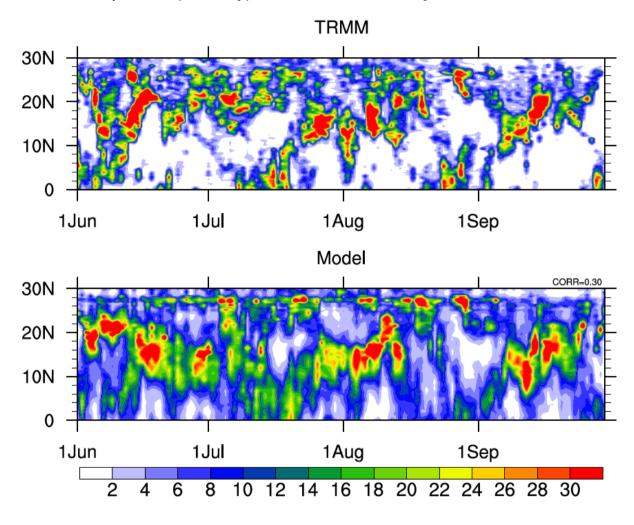
Precipitation (mm/day), Observation and Day-2 Forecasts, 83-93E

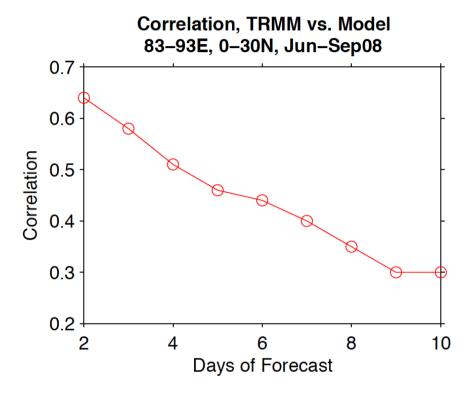


Precipitation (mm/day), Observation and Day-6 Forecasts, 83-93E



Precipitation (mm/day), Observation and Day-10 Forecasts, 83-93E





Spatial Scale of Meridional Propagation of Convection

Morlet Wavelet:

$$\psi_0(\eta) = \pi^{-0.25} e^{i\omega_0 \eta} e^{-\eta^2/2}$$

$$= \sqrt{\frac{2}{\pi}} e^{i\omega_0 \eta} \underbrace{\frac{1}{\sqrt{2\pi}} e^{-\eta^2/2}}_{\text{N(0,1)}}$$

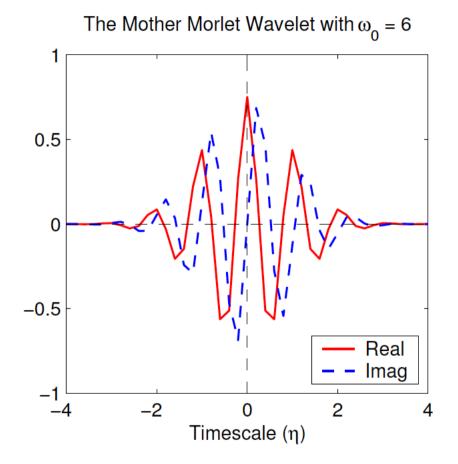
Wavelet transform of the time series is the convolution of a scaled and translated 98 version of this wavelet function:

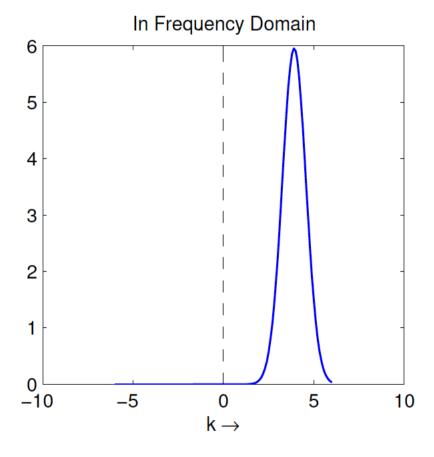
$$W_n(s) = \sum_{n'=0}^{N-1} y_{n'} * \psi^* \left(\frac{(n'-n)\delta t}{s} \right)$$

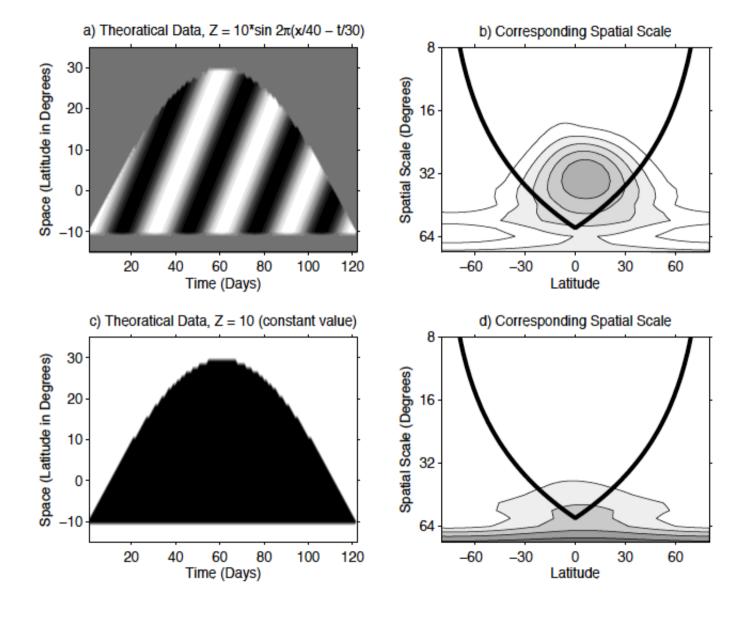
Final field after 2dimensional wavelet analysis:

$$w(s_y, y, s_t, t) = g(s_y, y, t) * \frac{1}{\sqrt{s_t}} \psi_0\left(\frac{t}{s_t}\right)$$

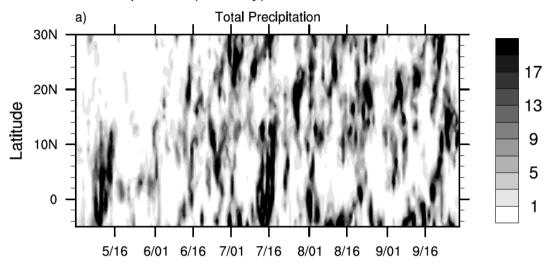
Chakraborty and Nanjundiah (2012, MWR)



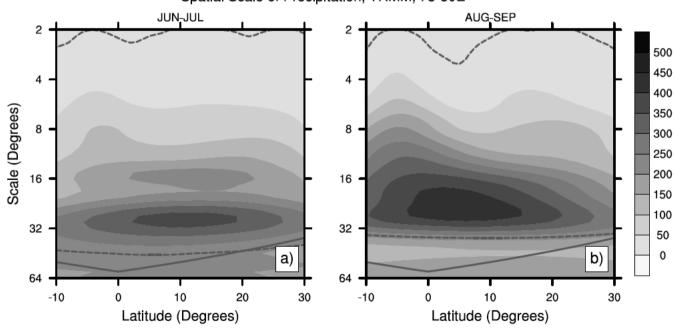


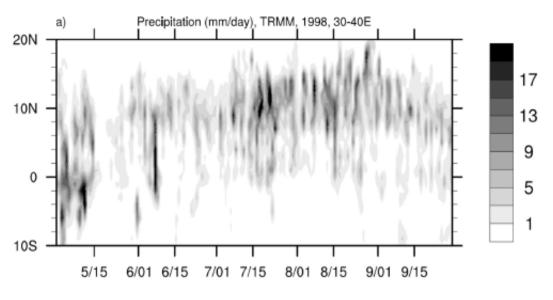


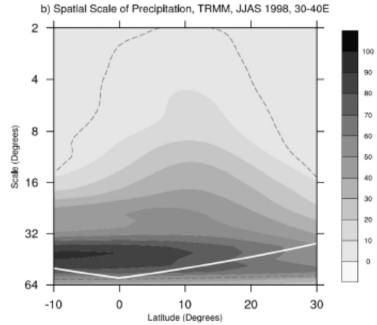
Precipitation (mm/day), TRMM, 1998, 75-80E

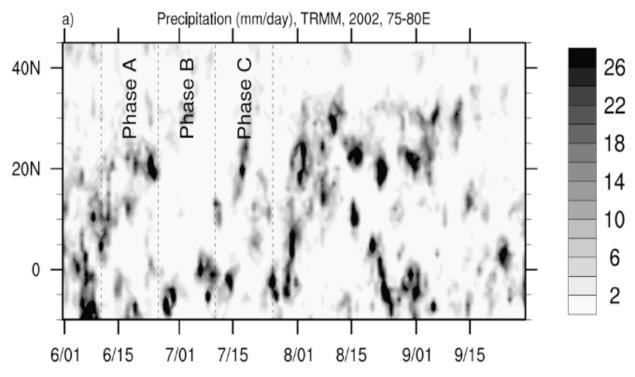


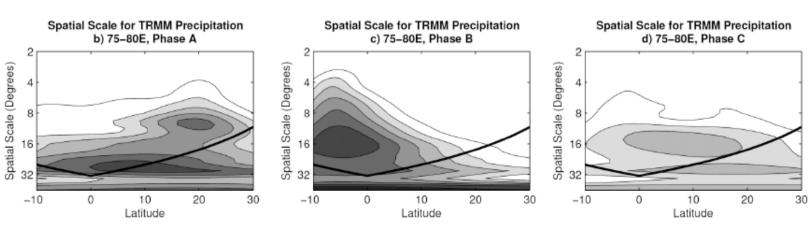
Spatial Scale of Precipitation, TRMM, 75-80E

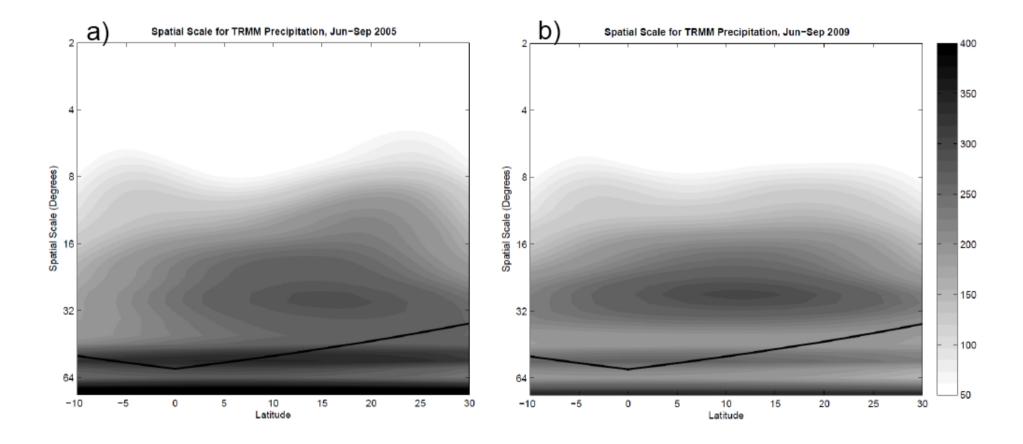




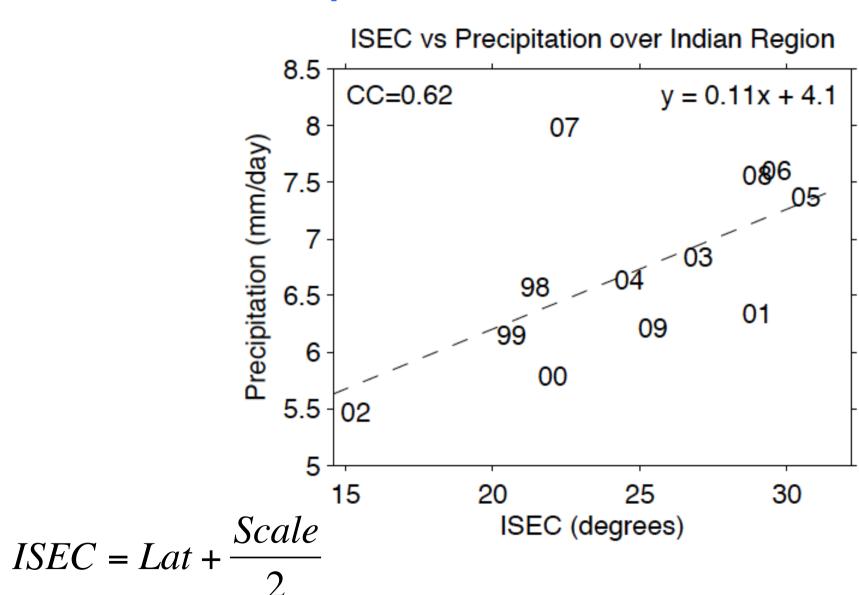


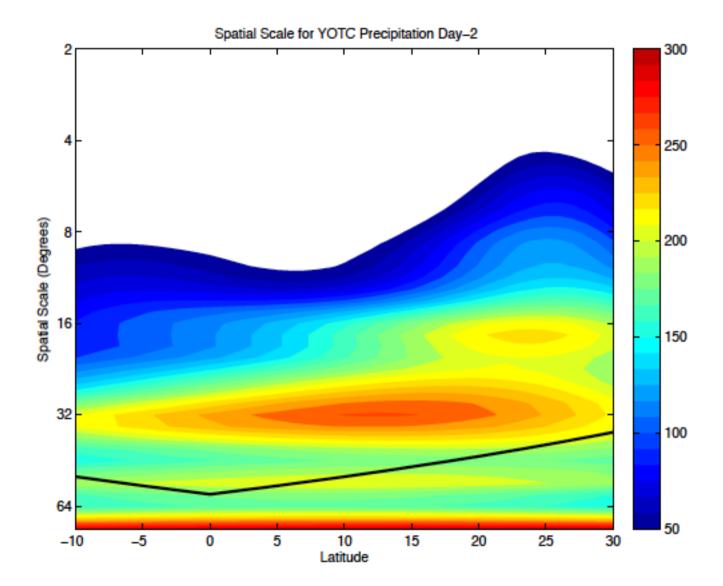


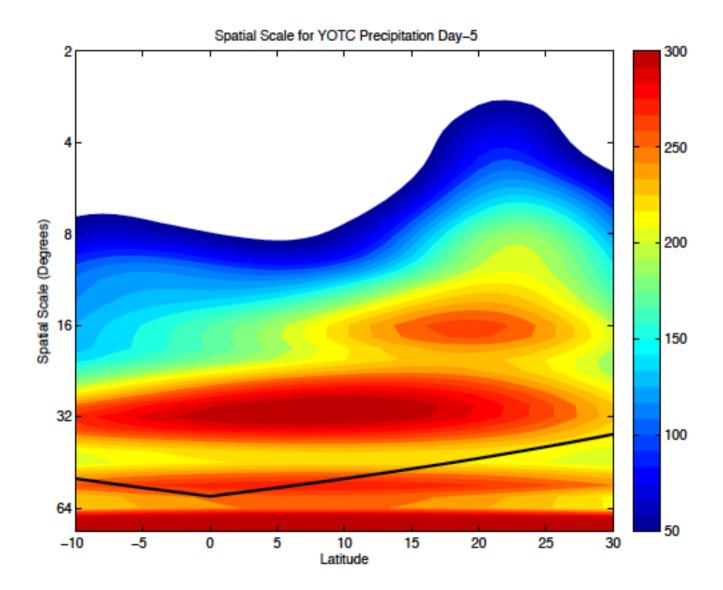


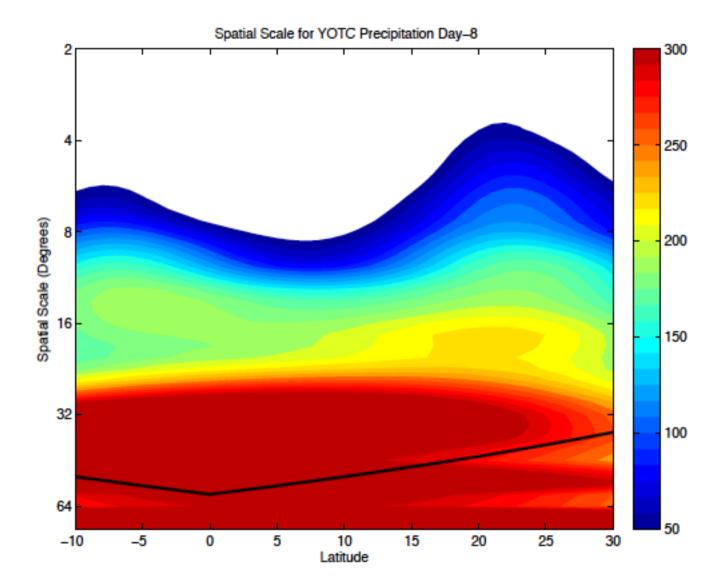


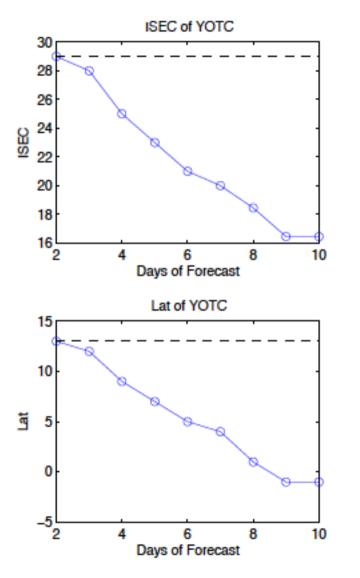
Index for the Spatial Extent of Convection

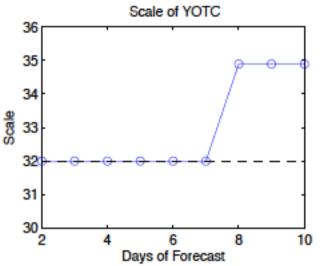






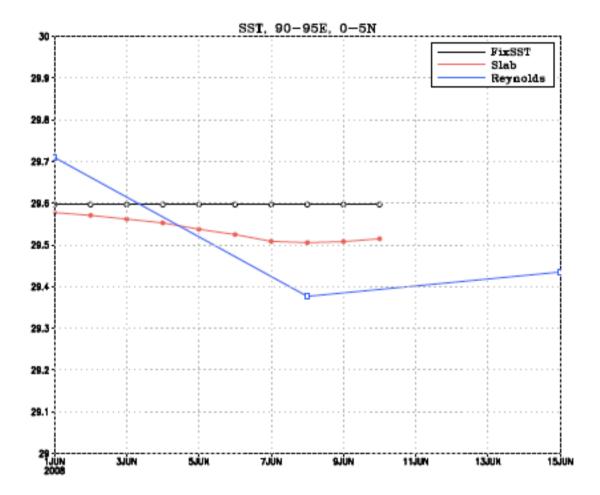




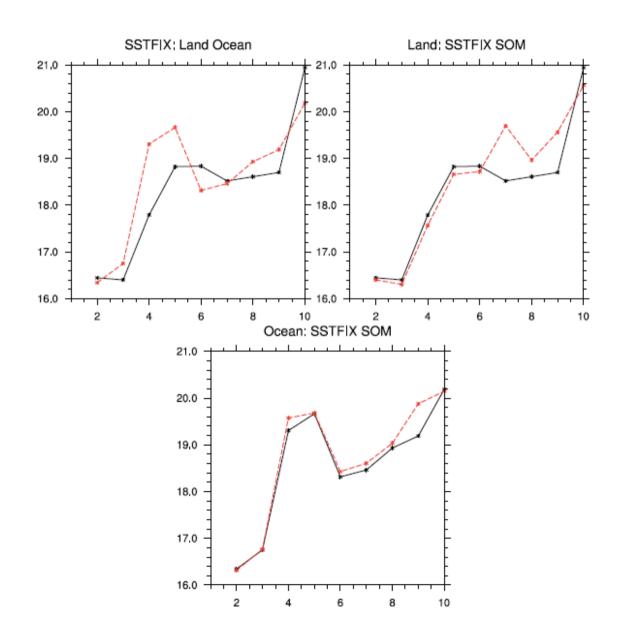


Model Details and Experiments

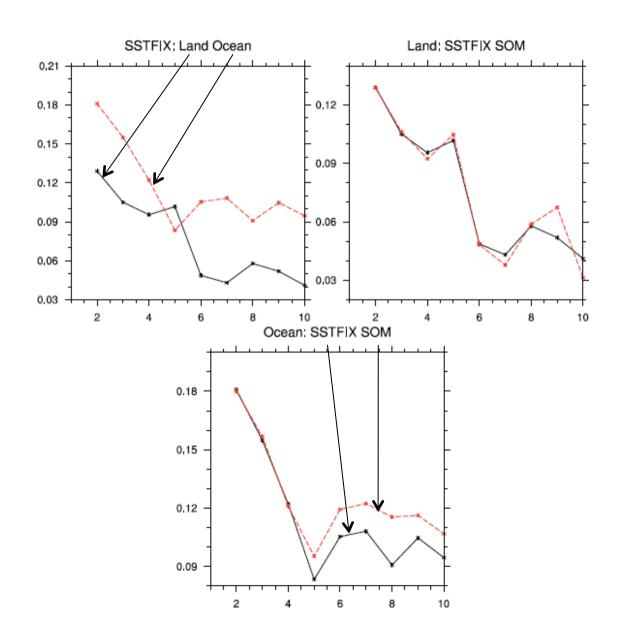
☐ Seasonal Forecast Model (SFM) with T170 resolution. ☐ Grid size is about 0.7 deg near the Equator. ☐ 10-day long forecast with fixed SST and slab ocean model. ☐ Slab ocean model (SOM) with climatological mixed layer depth, solar radiation transmission (based on Ohlmann and Siegel 2000) depends on clouds and angle of zenith angle (chlorophyll). ☐ 10-day long experiments were performed from end of May through June of 2008, starting at 00 GMT of every day.



RMS Error in Precipitation over South Asian Monsoon Region



Spatial Correlation in Precipitation over South Asian Monsoon Region



- 1) ECMWF model generally overestimates precipitation over tropical land and ocean for days 2 through 10 of forecasts.
- 2) Over the south Asian region, forecasted precipitation over the ocean is higher for longer lead time.
- 3) Rate of growth of error is larger over ocean as compared to that over land. This is probably due to lack of two-way air-sea interaction in the model.
- 4) Preliminary experiments suggest that at longer lead time skill over ocean can be improved with a slab ocean model coupled to the GCM.
- 5) An atmospheric model coupled with a slab ocean model seems to improve the medium range forecast skill.

A typical diagnostic cloud scheme

- o Radiation code of the model allows cloud to form at any layer.
- o Cloud amount is determined from the relative humidity.
- o If layer mean (mass weighted) relative humidity exceeds a specified threshold value cloud is present in that grid at that layer.
- The cloud amount is

$$cc = \left(\frac{rh - rh_c}{1 - rh_c}\right)^2$$

where rh_c is the threshold relative humidity:

⇒ High cloud amount is corrected using the convective precipitation (if non-zero):

$$cc_h = 4(0.247 + 0.126 \ln P - 0.3)^2$$

where P is precipitation in mm.

Cloud type	<i>rh</i> _c
High	0.40
Middle	0.45
Low	0.70

- Earlier cloud schemes used relative humidity (RH) threshold criteria for estimating cloud amount.
- Modifications to those schemes used other parameters like stability and vertical velocity as modifiers to RH based cloud fraction.

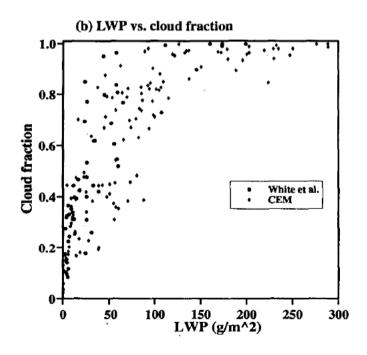
Prognostic Cloud Scheme

- ☐ Clouds and precipitation are tightly coupled to each other in nature.
- ☐ However, in large-scale global general circulation models (GCM), clouds are treated separately with precipitation.
- ☐ A prognostic cloud scheme includes and predicts cloud fractions and cloud liquid/ice water mixing ratio as prognostic variables.
- ☐ The cloud optical properties are interactive in such schemes.
- Moreover, prognostic cloud schemes can contain all important cloud microphysical processes such as autoconversion, accretion and cloud/rain water evaporation.
- Mass balance equation for cloud-water/ice content:

$$\frac{d}{dt}\int \rho_w dV = S_w - D_w$$

Cloud Scheme in GFS

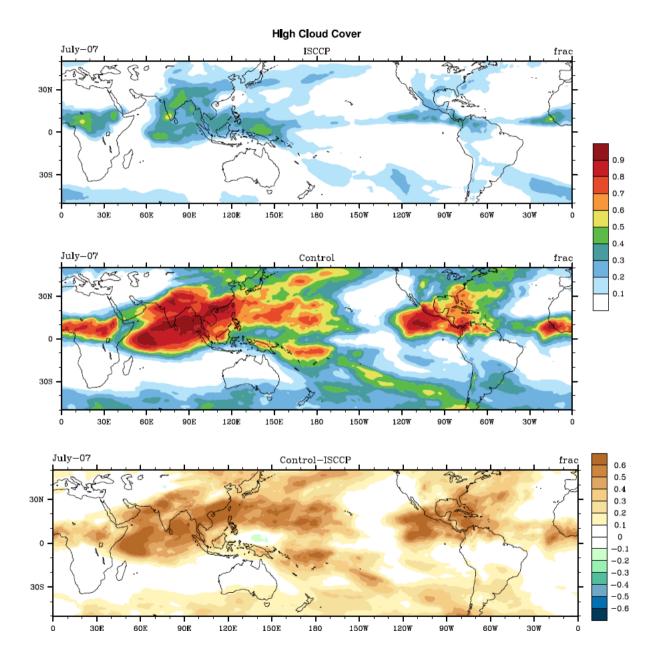
- The new version (v2) of the global forecast system (GFS) model uses a prognostic cloud scheme.
- The scheme is based on Zhao and Carr (1997), Xu and Randall (1996), and Moorthi et al (2001).
- Cloud fraction is calculated using cloud liquid and ice water mixing ratio $(q_i + q_i)$ and relative humidity (RH).

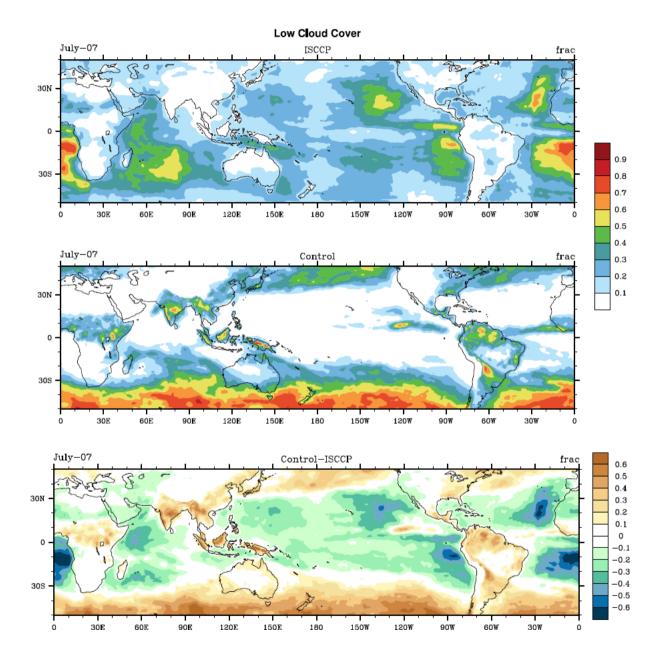


$$CC = RH^{0.25}(1 - \exp(-\frac{2000(q_l + q_i)}{((1 - RH)q_s)^{0.25}}))$$

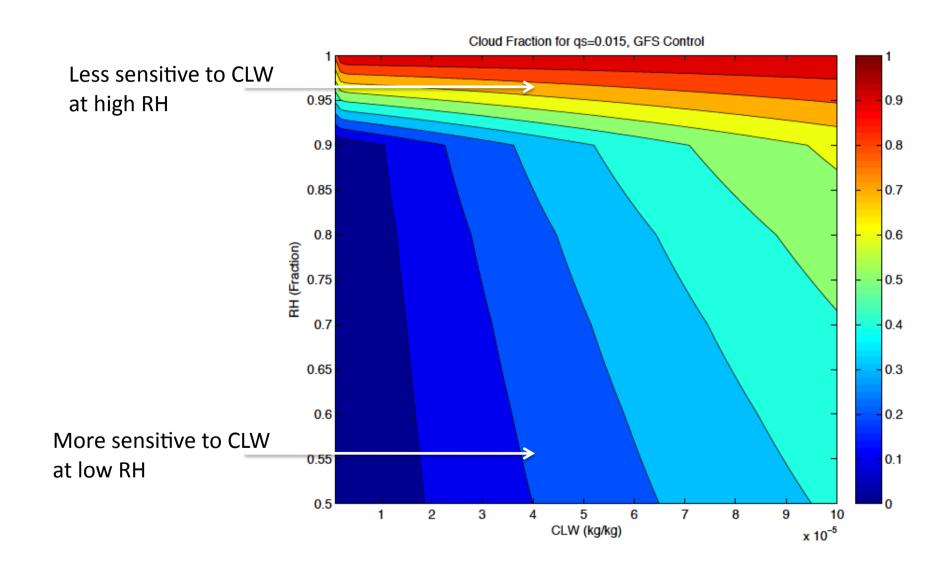
Model and Experimental Setup

- 1. Global Forecast System (GFS) model was used.
- 2. Resolution: T190 which corresponds to 0.64x0.64 degree in horizontal.
- 3. Model has 64 vertical hybrid sigma-pressure level.
- 4. Model was integrated for 35 days from 27 June 2007.
- 5. Initial condition from NCEP final analysis (FNL).
- 6. SST boundary condition prescribed from Reynolds weekly mean data.



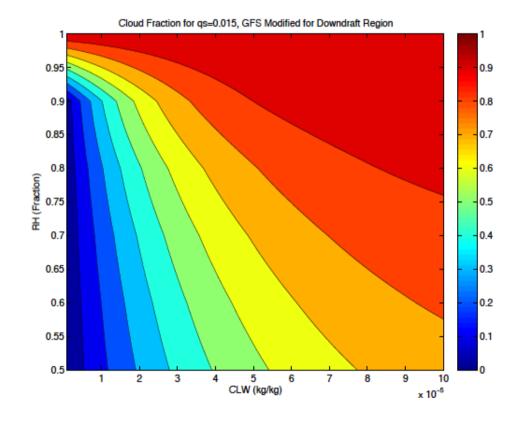


Synthetic Cloud Fraction in GFS Control

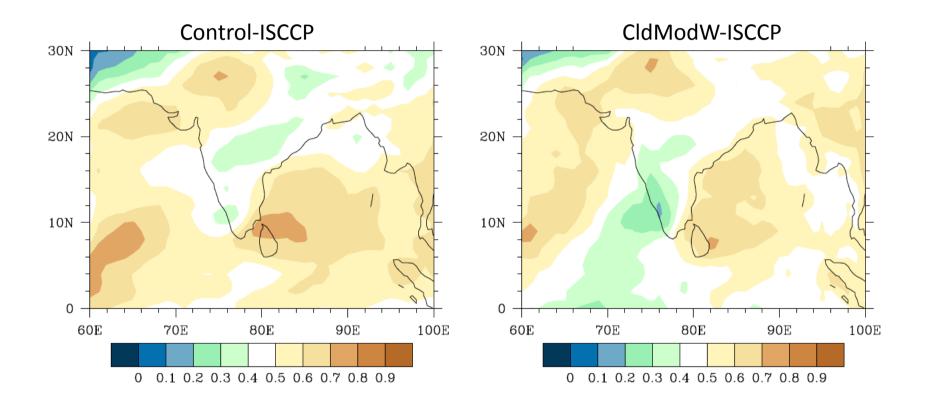


Synthetic cloud fraction in GFS Modified Scheme

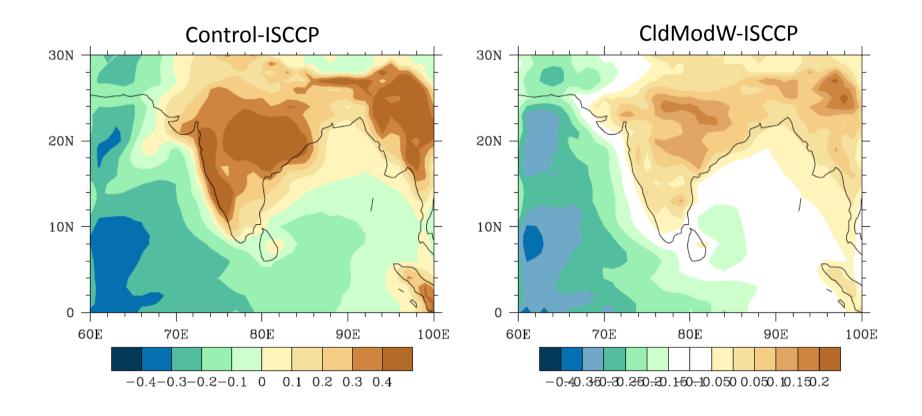
- 1. Modified in the cloud scheme for more sensitivity to CLW in low RH regime.
- 2. Introduced vertical velocity criteria for calculation of low cloud fraction.



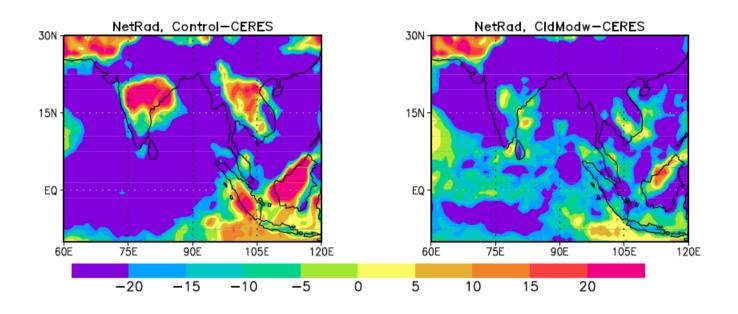
Error in High Cloud Cover, Jul07

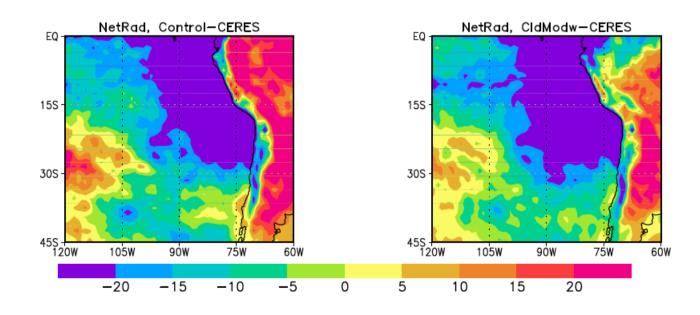


Error in Low Cloud Cover, Jul07

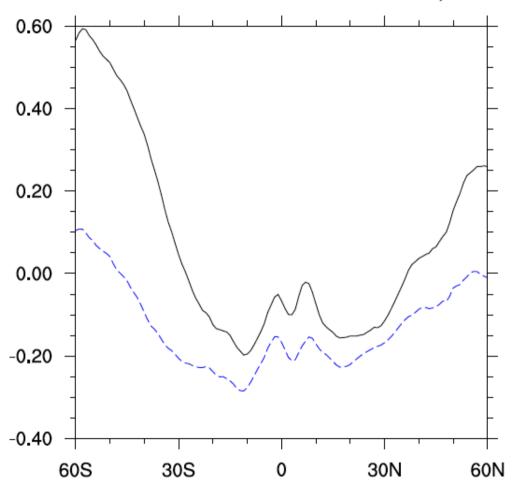


Error in TOA Net Radiation

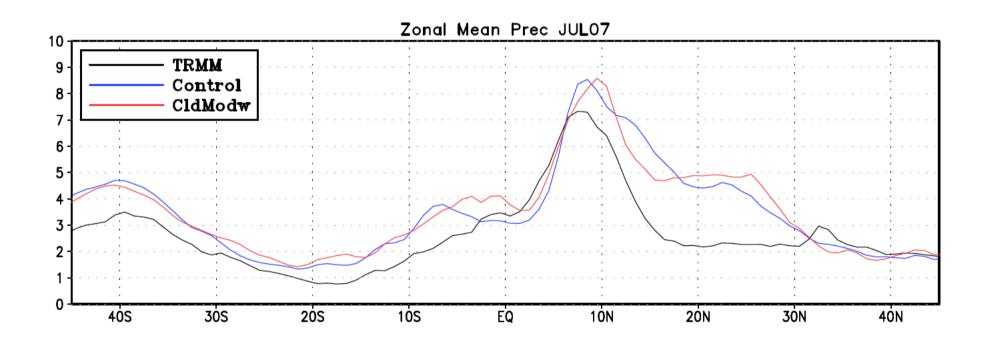




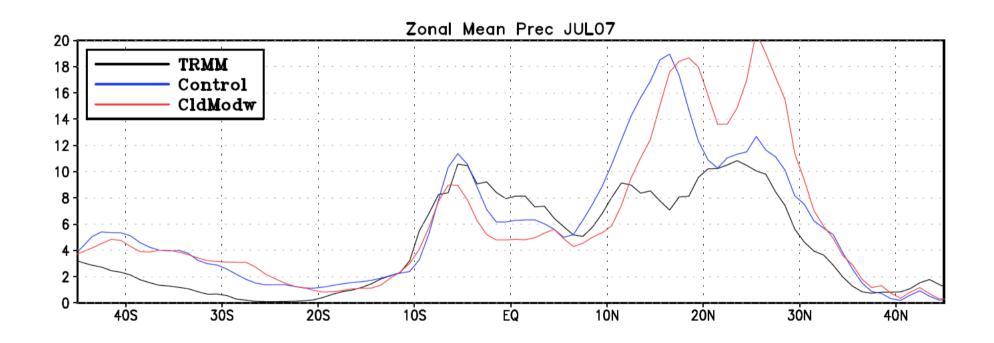
Error in Zonal Mean Low Cloud Cover, JUL07



Prec, 0-360E Mean



Prec, 70-100E Mean



Conclusions

- 1. Current GFS model has large error in simulating cloud covers.
- 2. These errors are propagated to error in radiation budget.
- 3. It is possible to correct some of those errors by modification in cloud parameterization scheme.
- 4. Such improvements can improve monsoon forecast by the model.
- 5. More observations in cloud microphysical scale required for better development of parameterization scheme.

